Assimilation of GPS Radio Occultation Data for Global Weather Prediction at CWB

> Jen-Cheng J. Chang CCU/DAS

### Outline

- GPS RO observations
- CWB's 3DVAR System
- 2D GPS Ray-tracing Operator
- Some testing results
- Summary
- Things to do

### What to assimilate?

**GPS Radio Occultation measurements**:

- 1. Excess phase: caused by the bending of the radio signal at two frequencies: 1227.6 MHz, 1575.4 MHz.
- 2. Excess Doppler frequency shift: estimated by the time derivative of excess phase.
- 3. <u>Bending angle</u> and impact parameter: derived from Doppler frequency shift based on satellite geometry (impact parameter is assumed constant at GPS and LEO).
- 4. *<u>Refractivity</u>*: calculated from bending angle through the Abel inversion (the refractivity is assumed spherically symmetric).
- 5. **Temperature and pressure:** retrieved from refractivity using the hydrostatic equation and neglecting water vapor content.

### Why bending angle? Accuracy

- The total effect of atmospheric refractivity along the ray path can be included.
- The effect of the ionosphere can be largely removed.
- Problems that are unique to GPS refractivity retrieval from bending angle can be avoided (e.g., the upper boundary condition for the Abel inversion and the ill-poseness of the Abel inversion under super-refraction).
- Providing a benchmark for developing a fast and accurate GPS refractivity assimilation method.
- Computational cost may be significantly reduced by running ray-tracing on multiple processors.

# Why not bending angle? *Efficiency*

### Why Refractivity

- 1. The computational cost is low to assimilate N.
- 2. A priori separation of temperature and moisture information is not required.
- 3. A weighted average (or a so-called *linearized non-local operator*) might be sufficient to account for the integrated effect of the atmosphere to CI

for the integrated effect of the atmosphere to GPS measurements.

# **3DVAR System at CWB**

### CWB's 3DVAR System

- Based on NCEP's SSI (version 1999)
- Operational since May 2003
- Official version: T179/L30 (i.e., 540 x 270 x 30), running with 3PE (on Fujitsu 5000)
- Testing version: T79/L30 (i.e., 240 x 120 x 30), running with 1PE
- Incremental approach: only 1 outer loop, with 100 inner loops (*currently testing 2 outer* updates with 70/30 iterations, respectively)
- No 3- and 9-hr forecasts for temporal interpolation to observational time.

### CWB 3DVAR (Contd.)

Analysis variables:

vorticity ( $\zeta$ ), unbalanced divergence(D'), unbalanced virtual temperature ( $T_v$ '), unbalanced log of surface pressure ( $\ln p_s$ '), specific humidity (q)

- Implicitly including a linear balance constraint
- Additional constraint: divergence tendency
- Background term at spectral space, observational terms at physical space

### Formula of CWB/3DVAR (i.e. NCEP/SSI) (Parrish and Derber, 1992)

#### Cost-function to be minimized:

$$J(\mathbf{w}) = \frac{1}{2}\mathbf{w}^T\mathbf{w} + \frac{1}{2}[\mathbf{y} - H(\mathbf{x}_b + \mathbf{C}\mathbf{w})]^T\mathbf{R}^{-1}[\mathbf{y} - H(\mathbf{x}_b + \mathbf{C}\mathbf{w})] + J_c$$

#### where

 $\mathbf{w} = \mathbf{C}^{-1}(\mathbf{x} - \mathbf{x}_b)$  Coefficients of error weighted analysis increments

 $\mathbf{x} = \mathbf{x}_b + \mathbf{C}\mathbf{w}$ 

 $\mathbf{B} = \mathbf{C}\mathbf{C}^T$  $\mathbf{x}_b$  $\mathbf{R} = \mathbf{F} + \mathbf{O}$ 

H

Analysis variables Background error covariance matrix 6-hr forecast of analysis variables Observational & Representative error covariance matrix (Nonlinear) observational (forward) operator Observations

### Formula (Contd.)

Gradient:

$$\frac{\partial J}{\partial \mathbf{w}} = \mathbf{w} - \mathbf{C}^T \mathbf{L}^T \mathbf{R}^{-1} [\mathbf{y} - H(\mathbf{x}_b + \mathbf{C}\mathbf{w})] = -\mathbf{f}(\mathbf{x})$$

where

W

 $\mathbf{y}_m$ 

$$\mathbf{L} = \frac{\partial H}{\partial \mathbf{x}} \text{ tangent linear operator of } H; \quad \mathbf{L}^T \text{ adjoint operator of } \mathbf{L}$$

$$\mathbf{Outer Loop:} \quad \text{At } m\text{-th iteration:} \qquad \mathbf{w} = \mathbf{w}_m \quad \mathbf{x} = \mathbf{x}_m \quad \mathbf{w}_0 = 0$$

$$\text{At } (m+1)\text{-th iteration:} \qquad \mathbf{w}_{m+1} = \mathbf{w}_m + \mathbf{d}$$

$$-\frac{\partial J}{\partial \mathbf{w}}\Big|_{m+1} = \mathbf{C}^T \mathbf{L}_m^T \mathbf{R}^{-1} \mathbf{y}_m - \mathbf{w}_m - (\mathbf{I} + \mathbf{C}^T \mathbf{L}_m^T \mathbf{R}^{-1} \mathbf{L}_m \mathbf{C}) \mathbf{d} = 0$$

$$\mathbf{here}, \qquad \mathbf{f}(\mathbf{x}_m) \qquad \mathbf{forcing vector} \qquad \mathbf{A}(\mathbf{x}_m)$$

$$\mathbf{coefficient matrix}$$

# Linear Conjugate Gradient Method

 $-\mathbf{p}_k^T \mathbf{d}_k$ 

or

 $\alpha_{k} =$ 

At the *k*-th update:

$$\mathbf{d}_k = \mathbf{d}_{k-1} + \boldsymbol{\alpha}_k \mathbf{p}_k$$

where,

$$\alpha_k$$
: step size

 $\mathbf{p}_k$  : search direction

Therefore,

$$\mathbf{f}_k = \mathbf{f}_{k-1} - \boldsymbol{\alpha}_k \mathbf{A} \mathbf{p}_k$$

$$\boldsymbol{\alpha}_{k} = \frac{\mathbf{p}_{k}^{T} \mathbf{f}_{k-1}}{\mathbf{p}_{k}^{T} \mathbf{A} \mathbf{p}_{k}}$$
to r

to minimize:  $\mathbf{f}_k^T \mathbf{A}^{-1} \mathbf{f}_k$ 

$$\frac{\mathbf{p}_{k}^{T} \mathbf{p}_{k}}{\mathbf{p}_{k}^{T} \mathbf{p}_{k}} + (\mathbf{L}_{m} \mathbf{C} \mathbf{p}_{k})^{T} \mathbf{R}^{-1} [\mathbf{y}_{m} - \mathbf{L}_{m} \mathbf{C} \mathbf{d}_{k-1}] - \mathbf{p}_{k}^{T} \mathbf{w}_{m}$$

$$\mathbf{p}_k = \mathbf{f}_{k-1} + \beta_{k-1} \mathbf{p}_{k-1}$$

 $\mathbf{p}_1 = \mathbf{f}_0$ 

$$\beta_k = \frac{\mathbf{f}_k^T \mathbf{f}_k}{\mathbf{f}_{k-1}^T \mathbf{f}_{k-1}}$$

(for orthogonality)

### 2D GPS Ray-tracing Operator



Courtesy of X. Zou



A slightly modified version of the 2D GPS ray-tracing operator from Zou et al. (1999) is implemented

### **Original Operator**

- Calculating N on the vertical velocity (half) level, but using variables (T and q) at the following model layer (full level) except p
- Calculating the geometric heights of vertical grids on the half-level, but treating *T* as given at the half-level in the hypsometric equation
- Results: a lower tropopause bias



### **Revised Operator**

- Calculating N on the model (full) layers, NOT the vertical velocity (half) level
- Calculating the geometric heights of vertical grids on the *full-layer*, and treating *T* back to where it belong

#### CWB Model's Vertical Grids



Figure 3.1: The finite difference vertical structure of the forecast model.

#### $L18 \rightarrow L30$

# Data Assimilation Procedure of GPS RO Observation



# Experiments

Name	Remarks
nogps	All other available data, except GPS (Step I only)
gpswt5	Including GPS observation, but with 10 <sup>5</sup> O-weighting
gpswt6	Including GPS observation, but with 10 <sup>6</sup> O-weighting

### Observational Weighting Profile Used



### Case Study

- July 4, 2002, 1200UTC
- From GFZ (GeoForschungs Zentrum) Potsdam CHAMP-ISDC (<u>http://isdc.gfz-</u> potsdam.de/champ/)
- 41 soundings during 09-15UTC

# **Observations Used**

Variables	Types	Amounts
Winds	rawinsonde < pibal < wind profile < NEXRAD < AIREP < ACARS < SATOB < SHIP < BUOY < surface SSM/I wind speed	111192
Temperature	rawinsonde 、AIREP 、SHIP	43088
Water Vapor	rawinsonde  SHIP	10567
Surface Pressure	rawinsonde  surface land  SHIP	13770
Bending Angle	GPS/RO	6586(41 soundings with vertical resolution 200m)

### **Data Distributions**









### **GPS** Soundings

GPS1 for 1 sounding test



### GPS\_only Exp (1 sounding)



- Systematically larger model bending angles:
  - lower  $p_{sfc}$ ,
  - dryer q

 $\rightarrow$ 

- warmer T<sub>v</sub>

### Analysis increments:









wgt5

- lower adjustment indeed.
- wgt6 : wgt5 ~ 5 : 1

### Analysis increments: q E-W cross section at lat = $47.2^{\circ}N$ & at $\sigma$ = 0.5658













### Analysis increments: T<sub>v</sub> E-W cross section at lat = 47.2 & at $\sigma$ = 0.5658

not as localized as q

•~6:1









wgt5

### Analysis increments: v E-W cross section at lat = 47.2 & at $\sigma$ = 0.5042

•~6:1









### Analysis increments: U N-S cross section at lon = 172.4E & at $\sigma$ = 0.5042

•~6:1









## **Multi-sounding Results**

### **Cost-Function**



gpswt5

# Analysis Increments (P<sub>sfc</sub>)

#### gps\_only (wt6)







gpswt5

# Analysis Increments (q)

#### norad





#### gps\_only (wt6)





#### gpswt5

# Analysis Increments $(T_v)$

#### gps\_only (wt6)







gpswt5

# Analysis Increments: u

#### gps\_only (wt6)







nogps

gpswt5

### Analysis Increments: v

#### gps\_only (wt6)







gpswt5

### Forecasts

### Analysis: Day 0



#### gpswt5 - nogps

#### nogps / gpswt5

### Forecasts: Day 3



#### gpswt5 - nogps

### Forecasts: Day 5



#### gpswt5 - nogps

#### nogps / gpswt5

### Anomaly Correlation (NH: 20°N-80°N / SH: 80°S-20°S)

	Exp	24-hr	48-hr	72-hr	96-hr	120-hr
SLP	gpswt5	0.8695 / 0.9290	0.8374 / 0.8965	0.7437 / 0.8520	0.6696 / 0.8255	0.6636 / 0.7552
	nogps	0.8695 / 0.9293	0.8380/ 0.8965	0.7433 / 0.8524	0.6683 / 0.8230	0.6630 / 0.7497
500 H	gpswt5	0.9380 / 0.9486	0.9380 / 0.9315	0.8831 / 0.9600	0.7873 / 0.8520	0.7590 / 0.7492
	nogps	0.9382 / 0.9487	0.9383 / 0.9309	0.8829 / 0.9555	0.7869 / 0.8495	0.7598 / 0.7413

(Yellow means better!)

### Root-Mean-Squared Errors (NH / SH)

	Exp	24 hr	48 hr	72 hr	96 hr	120 hr
Slp (mb)	gpswt5	3.0394 / 3.8771	3.0718 / 4.7778	<mark>3.8642</mark> / 6.0524	4.5628 / 6.7831	4.5594 / 7.9602
	nogps	3.0399 / 3.8722	3.0666 / 4.7803	3.8738 / 6.0488	4.5845 / 6.8404	4.5691 / 8.0626
500 H (m)	gpswt5	26.5011 / 37.9443	24.5412 / 43.8359	31.5442 / 52.9972	42.8493 / 65.8923	43.2711 / <mark>82.1605</mark>
	nogps	26.4472 / 37.9236	24.4784 / 44.0226	31.5891 / 53.1366	41.9354 / 66.4673	43.2277 / 83.5716
850 T (C)	gpswt5	1. <mark>8392</mark> / 2.7345	2.1729 / 3.0421	2.5124 / 3.5993	2.8975 / 4.0409	3.1161 / 4.3594
	nogps	1.8392 / 2.7281	2.1707 / 3.0409	2.5127 / 3.6115	2.8959 / 4.0756	3.1116 / 4.3951
200 Wind (m/s)	gpswt5	5.0369 / 6.6639	7.4152 / 8.2302	10.1056 / 10.0843	12.3924 / 13.1967	13.2515 / 15.9940
	nogps	5.0392 / 6.6706	7.4168 / 8.2561	10.1036 / 10.1231	12.3933 / 13.1900	13.2855 / 16.0811

### Summary

- A minorly revised 2D ray-tracing operator and its tangent-linear/adjoint operators (Chang et al., 2003, based on Zou et al., 1999) are currently implemented and tested on CWB/GFS.
- Though marginal, the forecasting impact in this case study is generally positive, which is encouraging!

# Summary (II)

• Upper-bound maximal analysis increments (GPS\_only\_wt6):

 $p_{sfc} \sim 4 hPa$   $T_v \sim 7 K$   $q \sim 5 g/kg$  $u, v \sim 2.5 m/s$ 

- GPS\_only 1-sounding tests suggest: maximal analysis increments are about 5-7 times smaller for the wt5 experiment.
- CWB's analysis increments without GPS:

p<sub>sfc</sub> ~ 1 hPa T<sub>v</sub> ~ 4-7 K q ~ 2-3 g/kg u, v ~ 10-15 m/s

# Summary (III)

- Additional increments added by GPS observation are generally one order smaller, except moisture (1-2 g/kg) and lower-level temperature (2 K).
- Substantial differences occurred in Day-5 forecasts.

## Things To Do

- Impact studies on CWB/GFS analysis and forecasts
- QC
- O matrix
- Speed up
  - Local refractivity operator
  - Data thinning
  - Parallelizing
- Linearized non-local refractivity operator