

Regional Climate Studies in East Asia

At Climate Dynamics Lab./ASNTU

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Department of Atmospheric Sciences

National Taiwan university

Part I: ISO vs. East Asian summer monsoon onset and TC

Part II: Interannual variability of summer rainfall

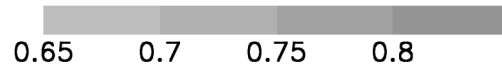
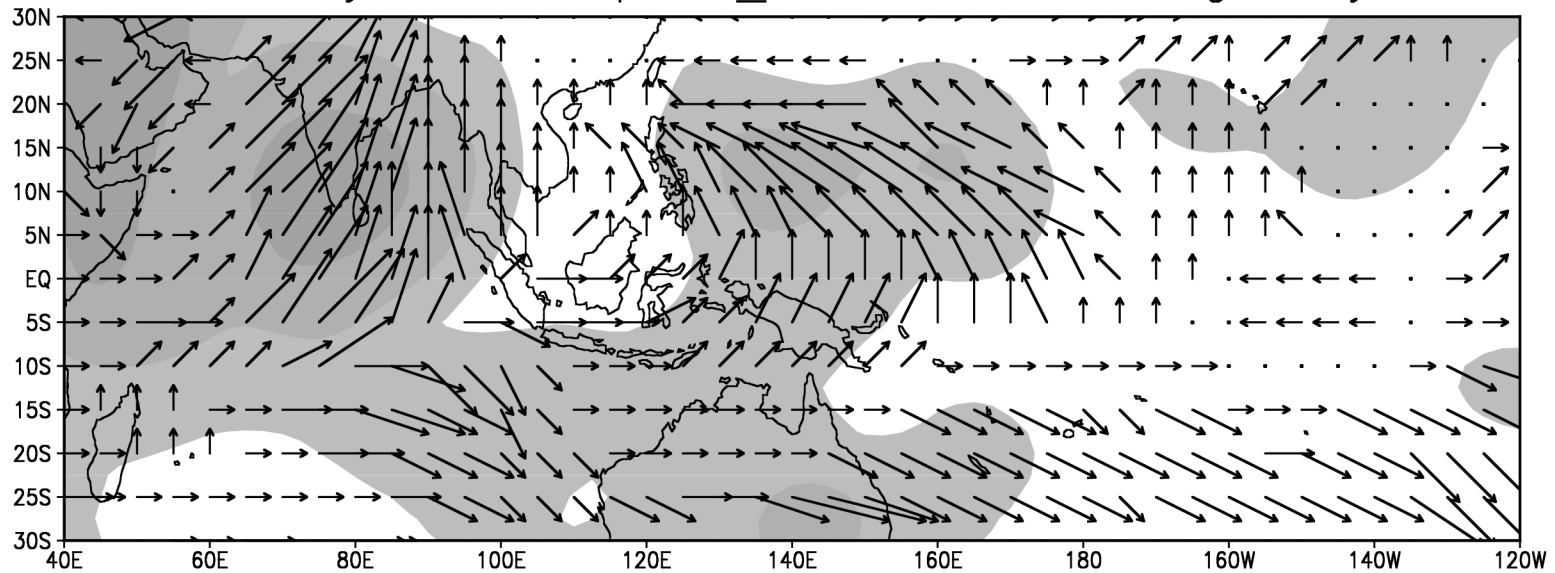
Part III: Regional Climate Simulation

Climate Dynamics Laboratory
<http://hsu.as.ntu.edu.tw>



Propagation tendency of ISO (April to October)

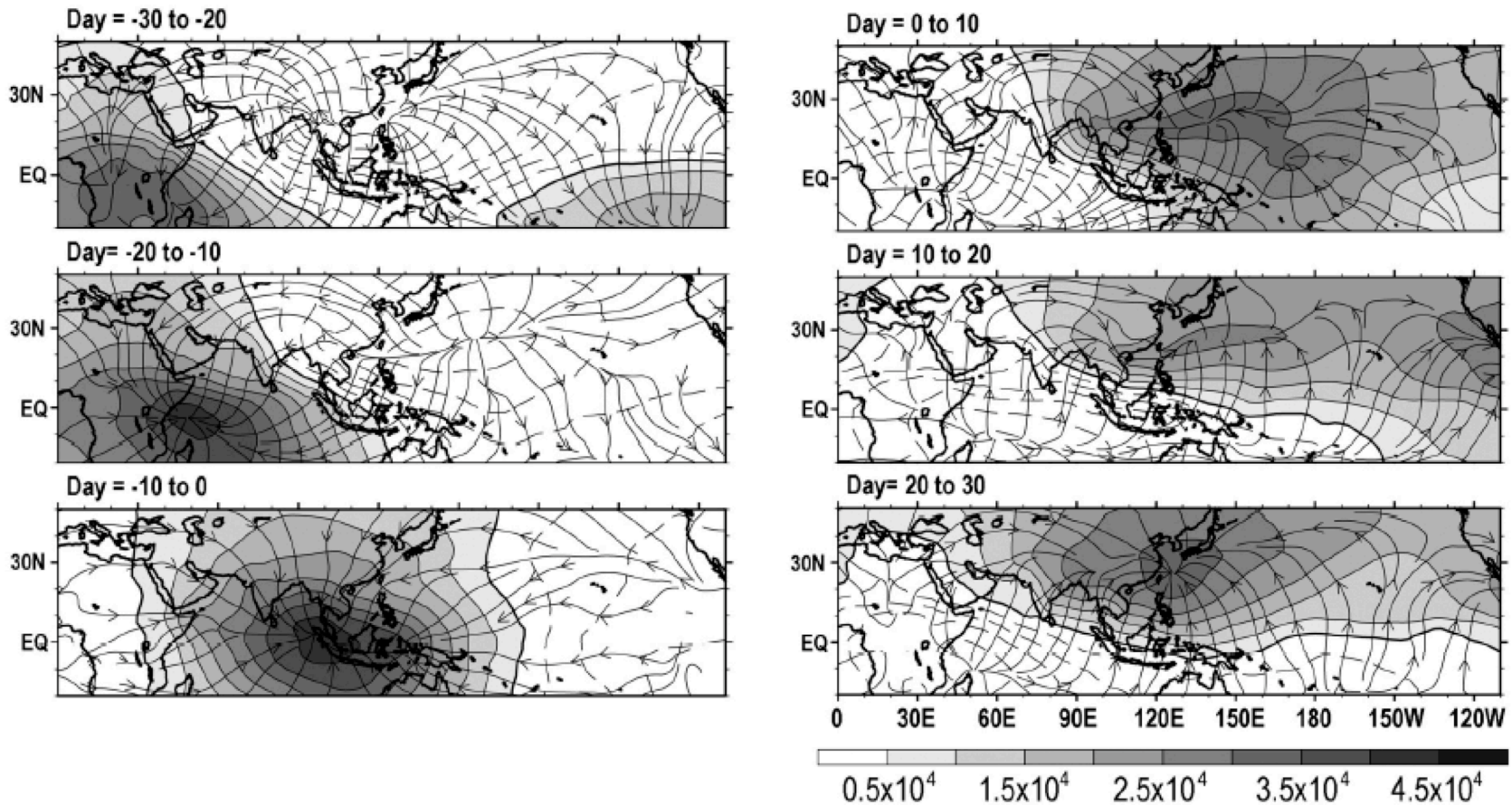
OLR 20–100dy-filtered spatial_filtered summer lag +5dy vectors



EA summer monsoon onset is strongly affected by MJO in some years

C.-W. Hung and H.-H. Hsu (2005)

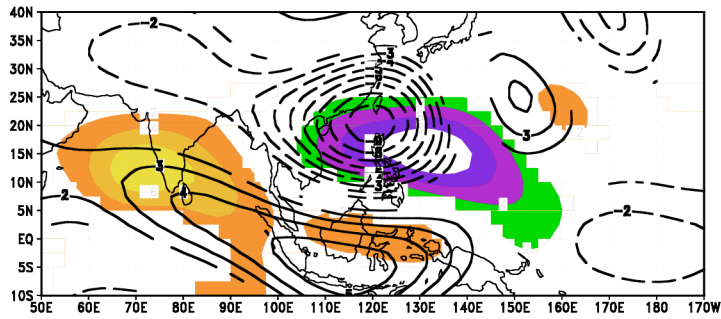
ERA40 850hPa Moisture divergence



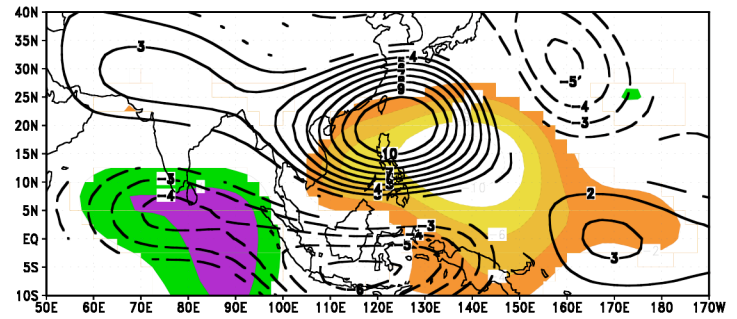
**Intraseasonal Oscillation
in the Western North Pacific
during Northern Summer**

H.-H. Hsu and C.-H. Weng (2001)

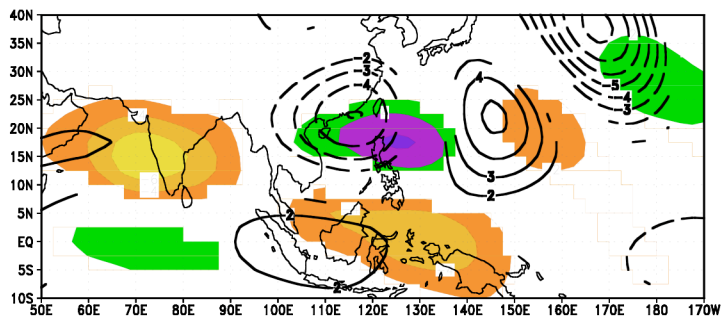
(a) vor850 - olr day -15



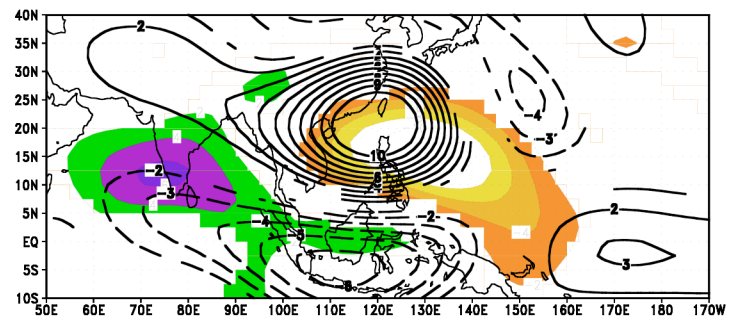
(d) day 0



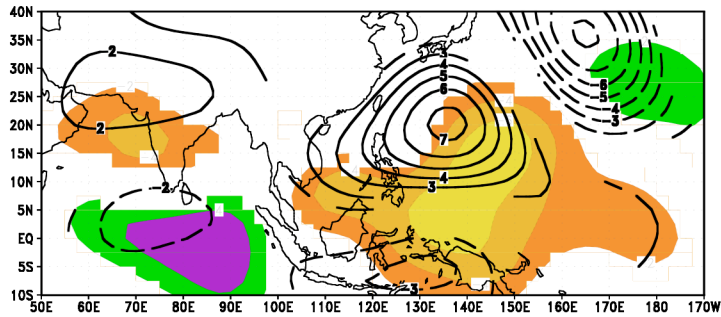
(b) day -10



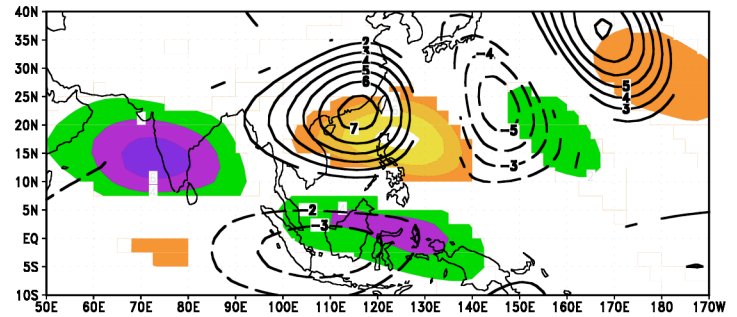
(e) day 5



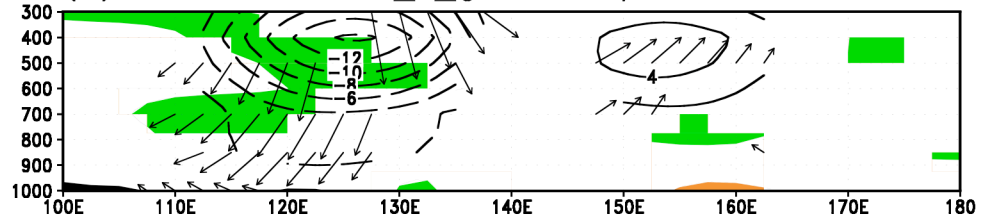
(c) day -5



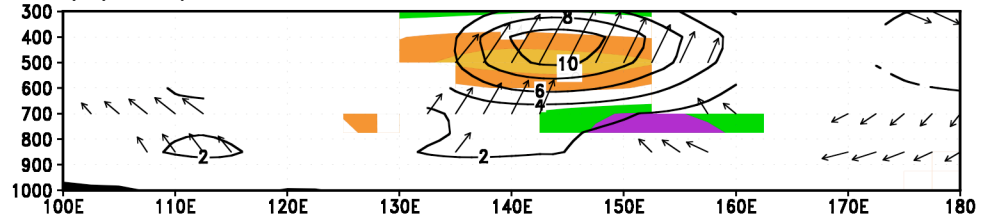
(f) day 10



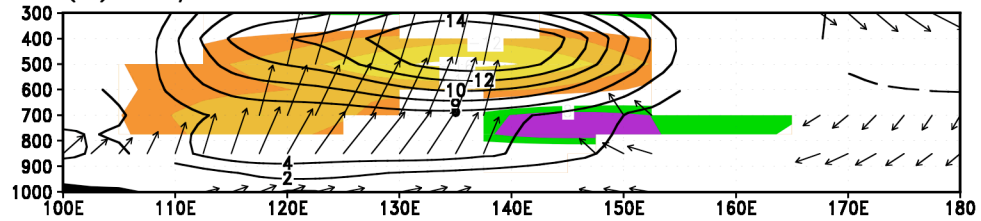
(a) dbh, uw, theta_e_grad day -10



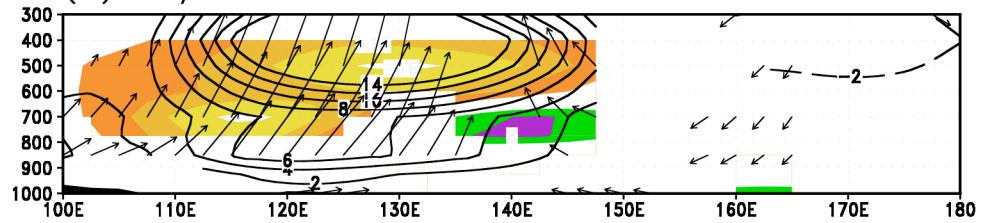
(b) day -5



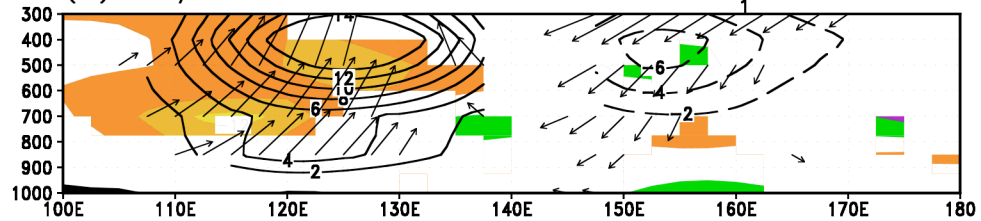
(c) day 0



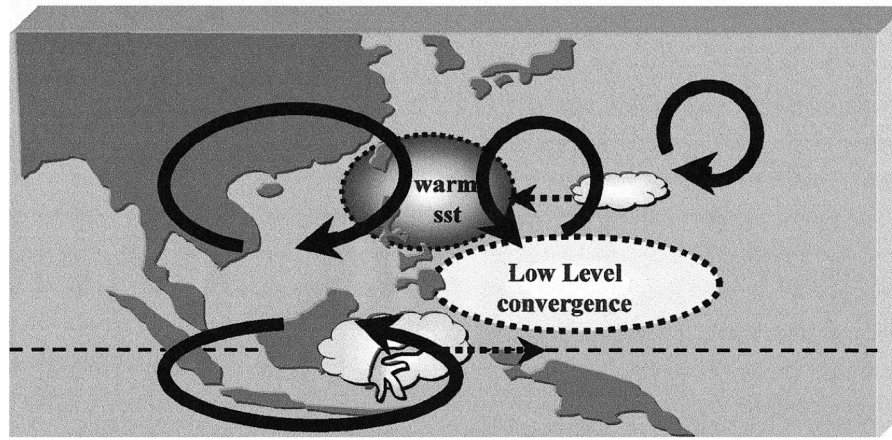
(d) day 5



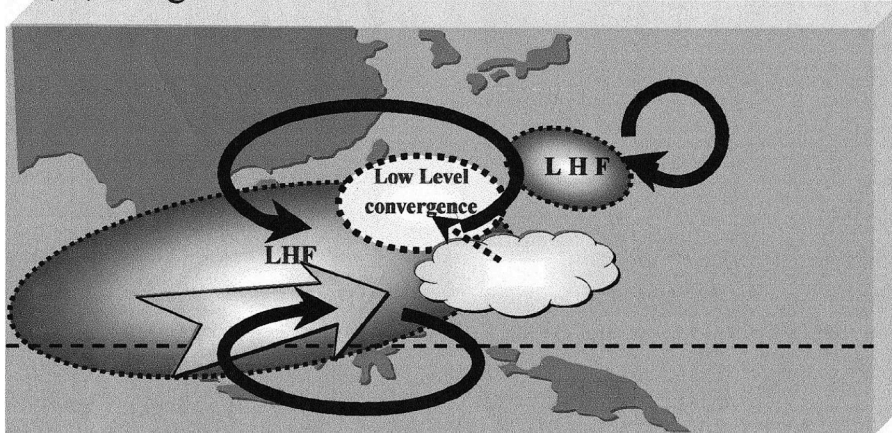
(e) day 10



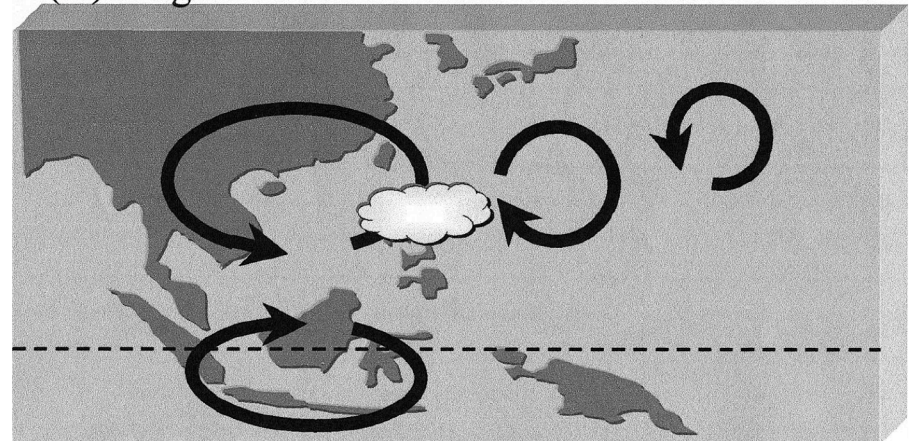
(a) Stage I



(b) Stage II



(c) Stage III



Strong ISO-TC Coupling in 2004 Typhoon Season

H.-H. Hsu and Y.-L. Chen (2005)

- recurring tracks - 10 typhoon landfalls in Japan
- appearing in clusters

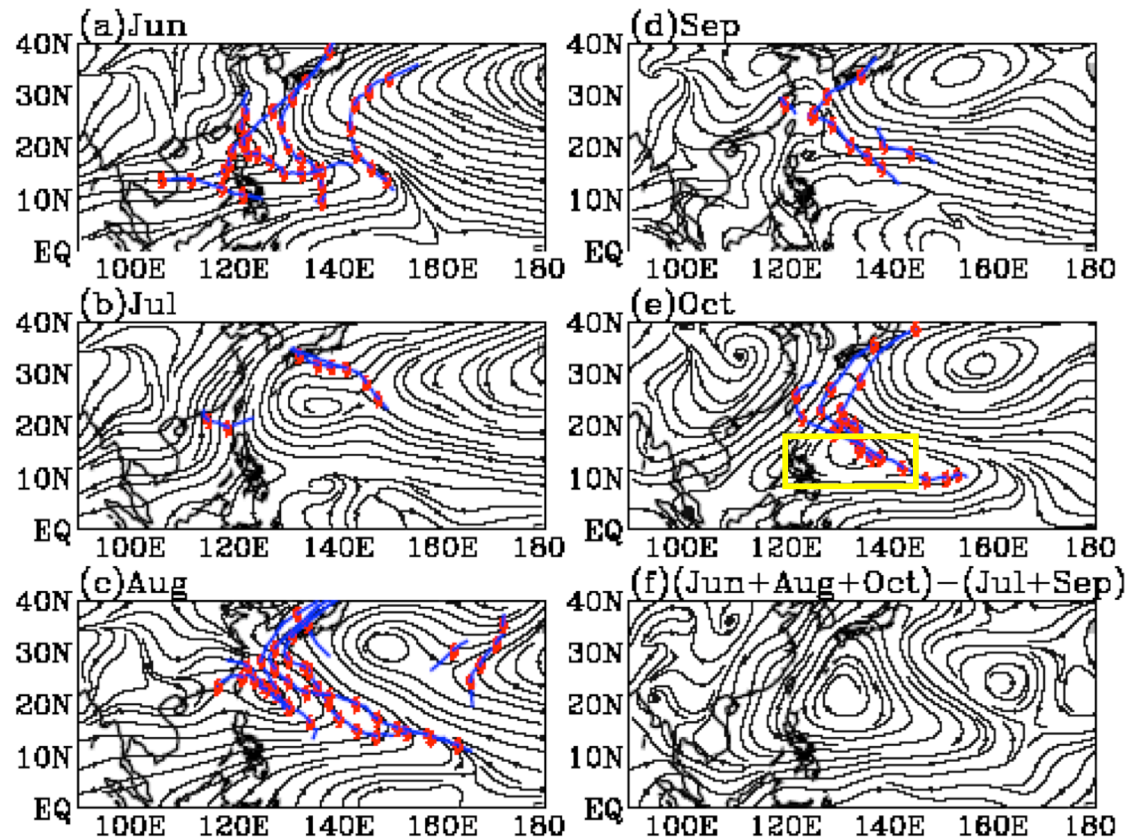
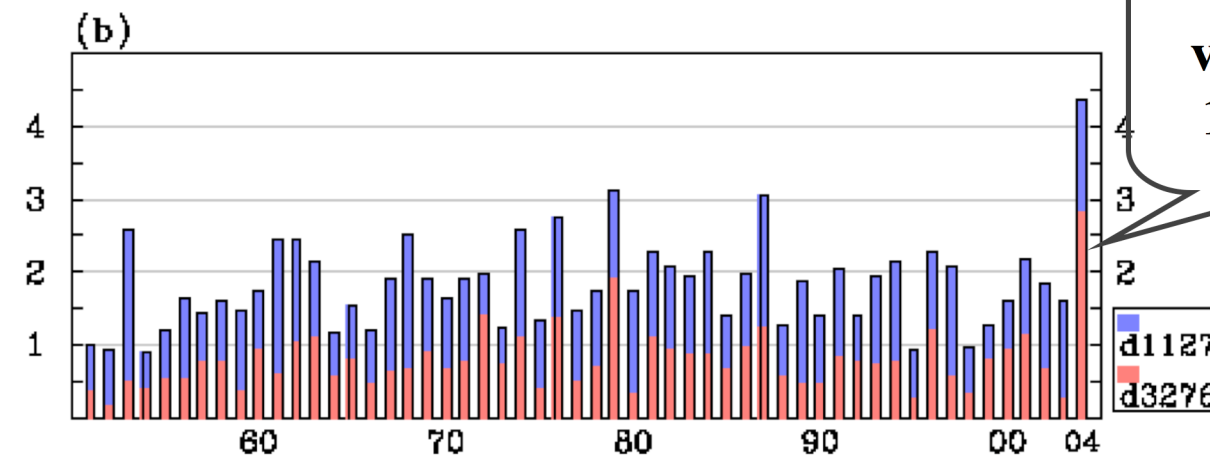
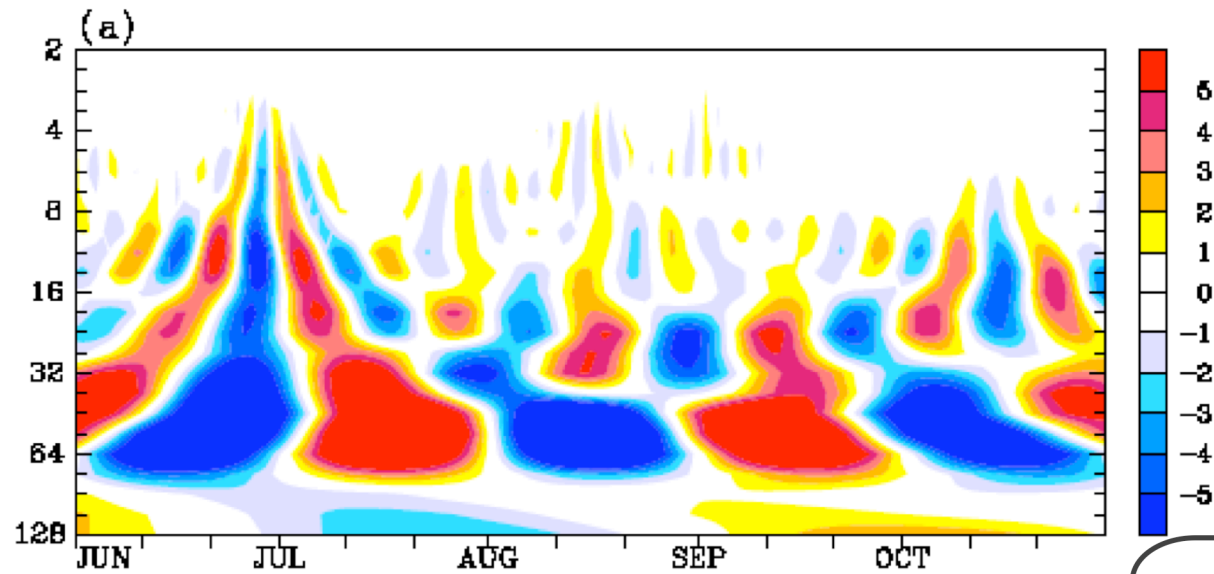


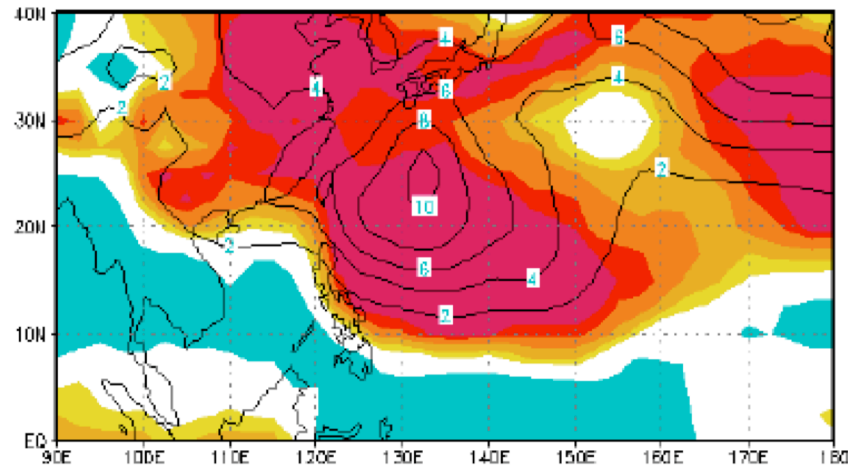
Figure 1. Monthly mean 850 hPa streamline and typhoon tracks from June to

monsoon trough index: MSLP (10N-20N, 120E-150E)

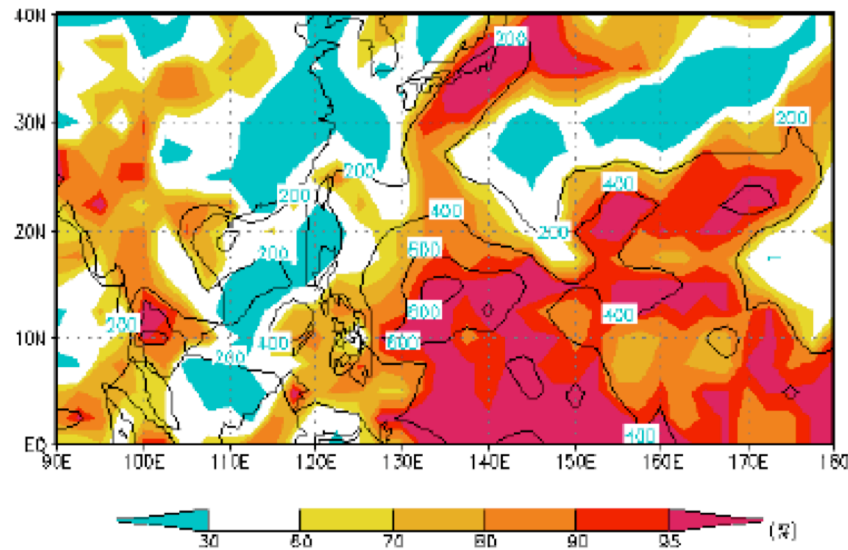


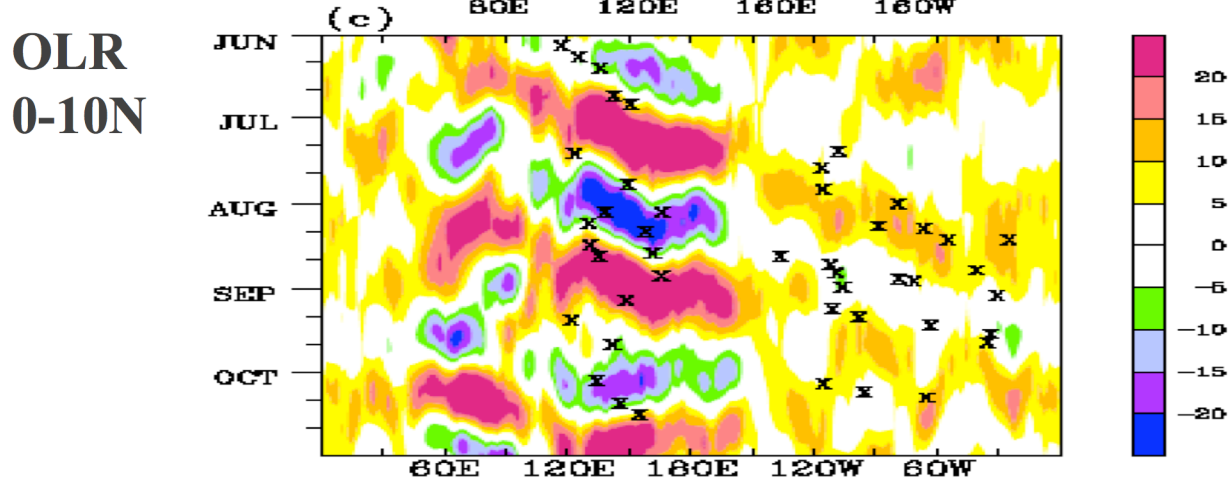
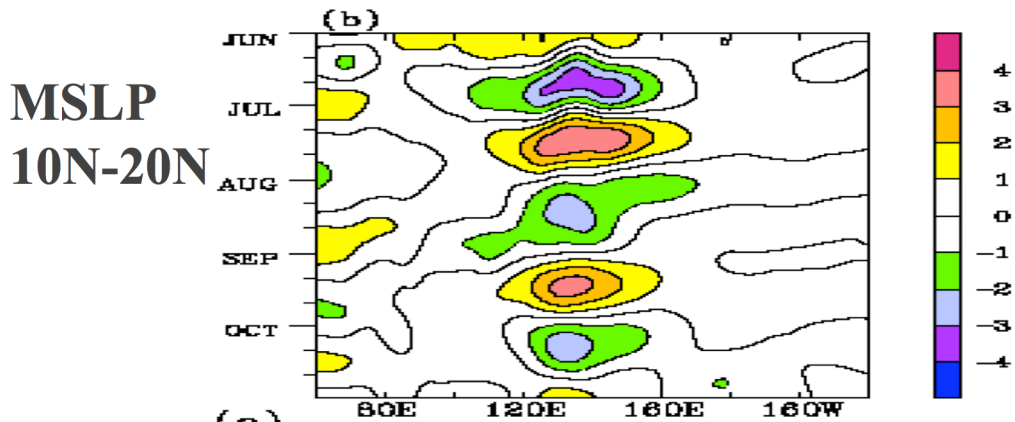
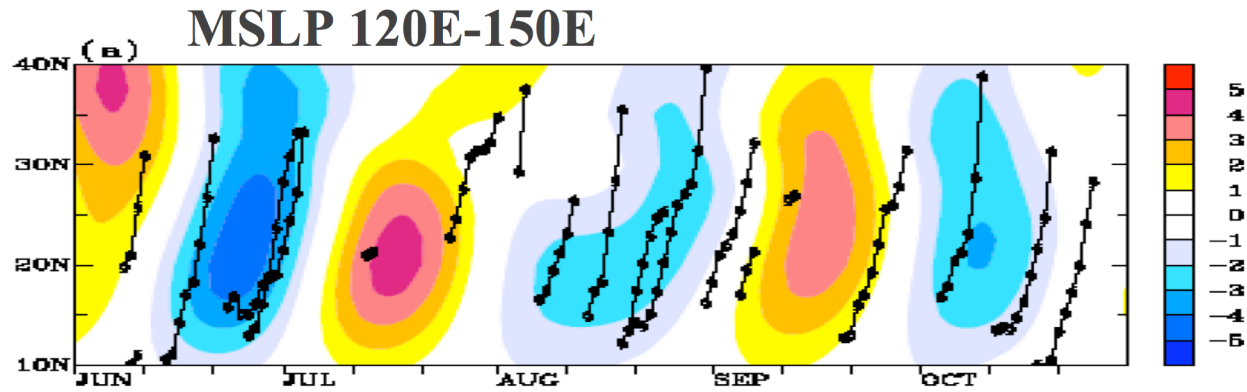
32-76-day variance and percentile in 2004

(a) MSLP



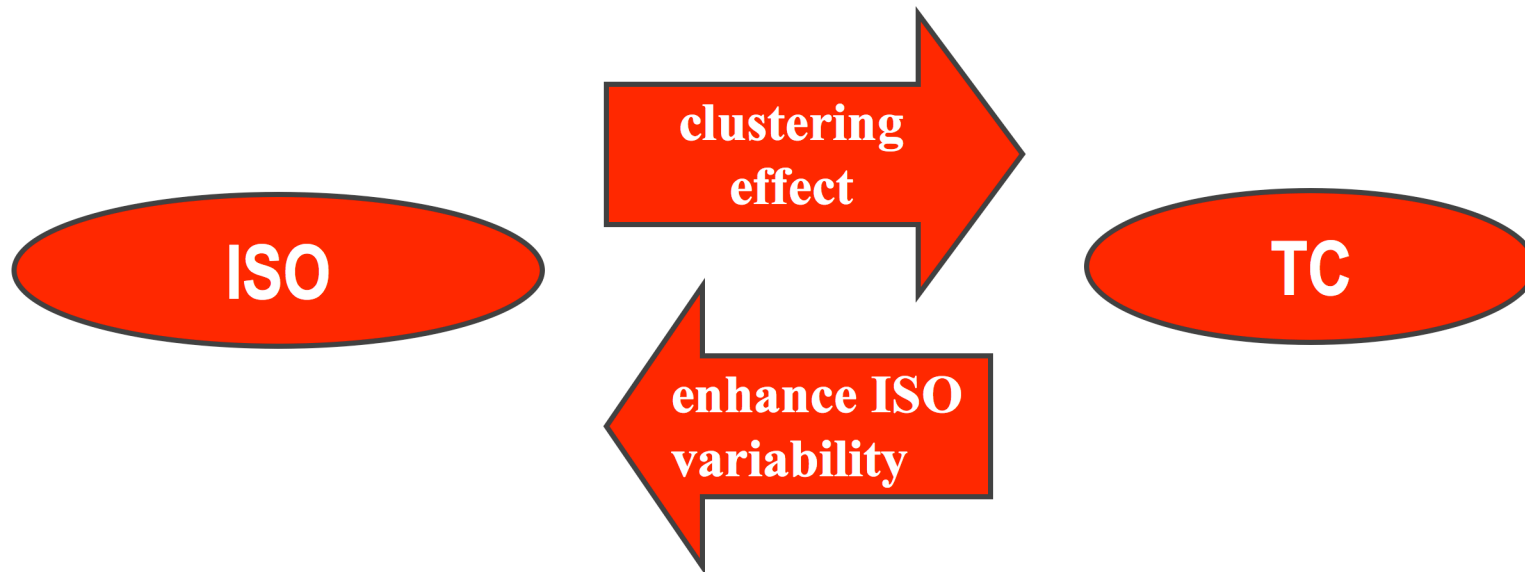
(b) OLR

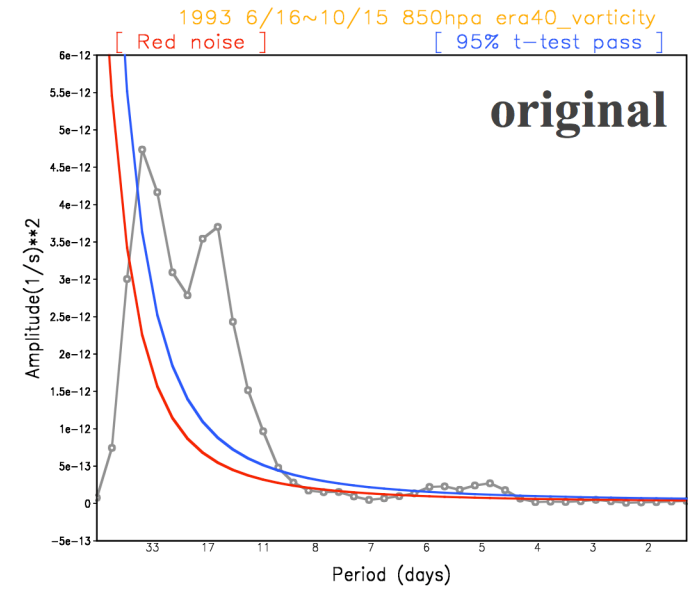




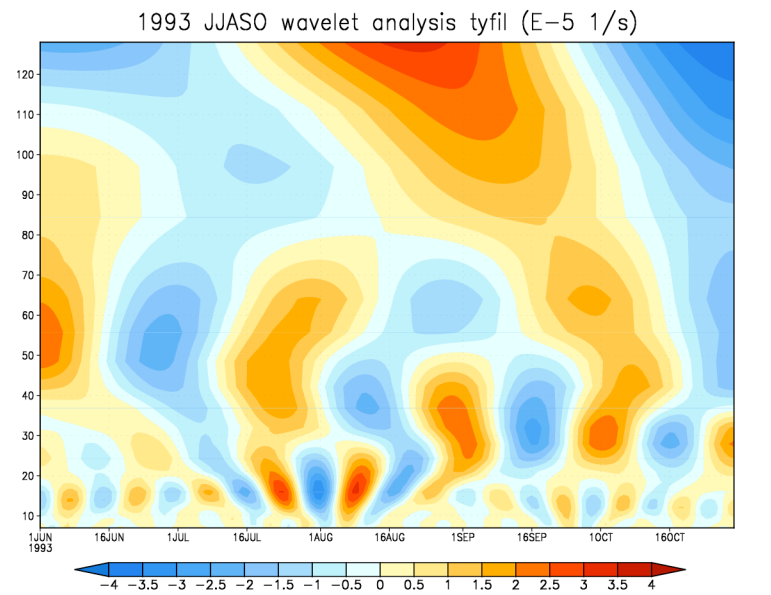
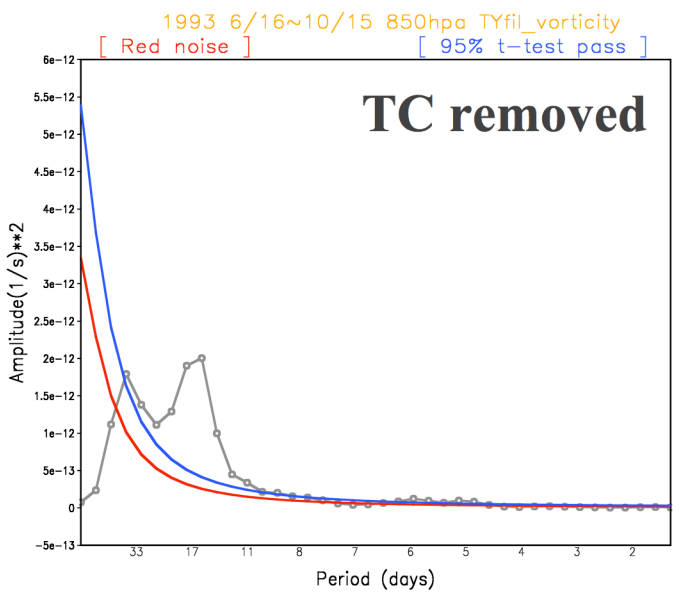
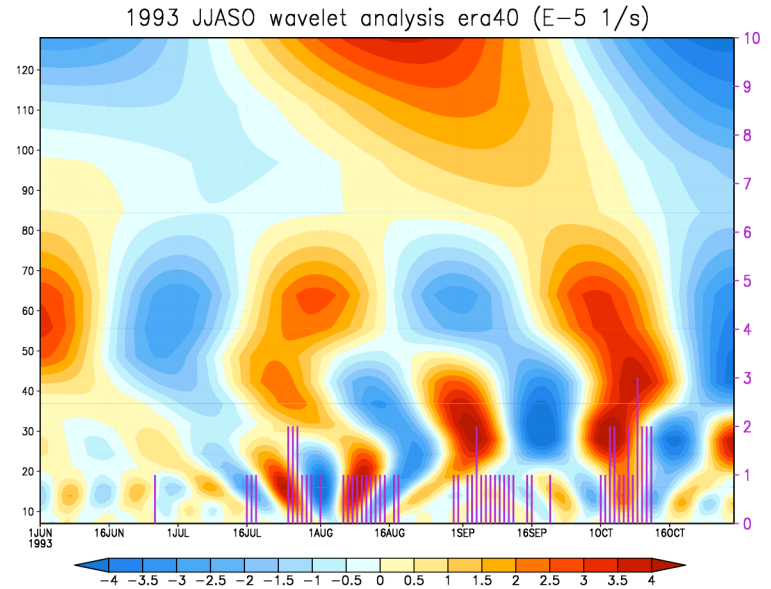
Interaction between ISO and TC

H.-H. Hsu, A.-K. Lo, and C.-C. Wu (2005)

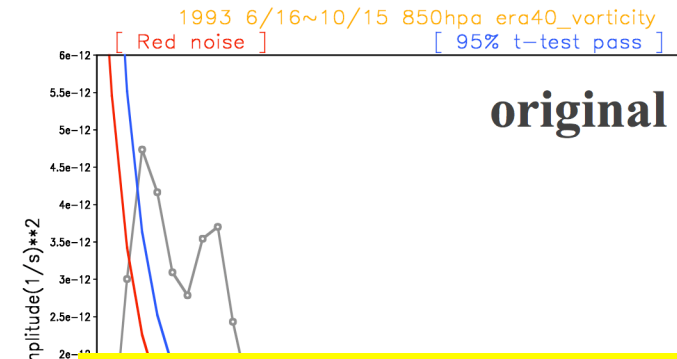




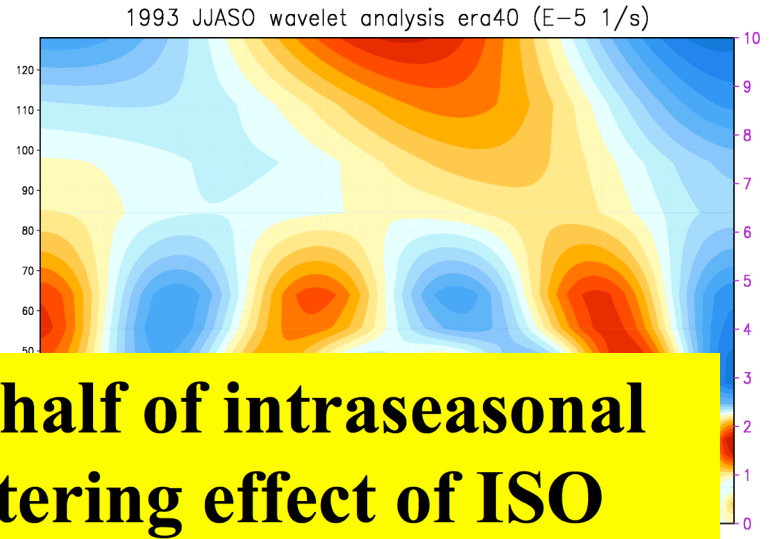
1993



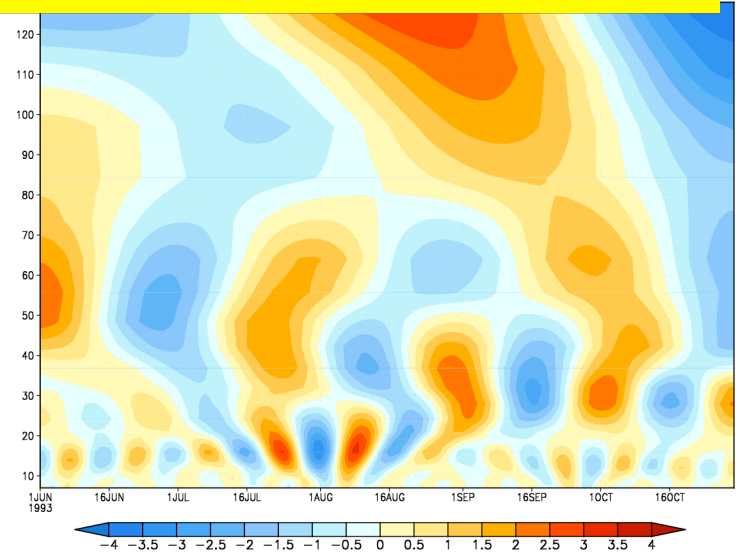
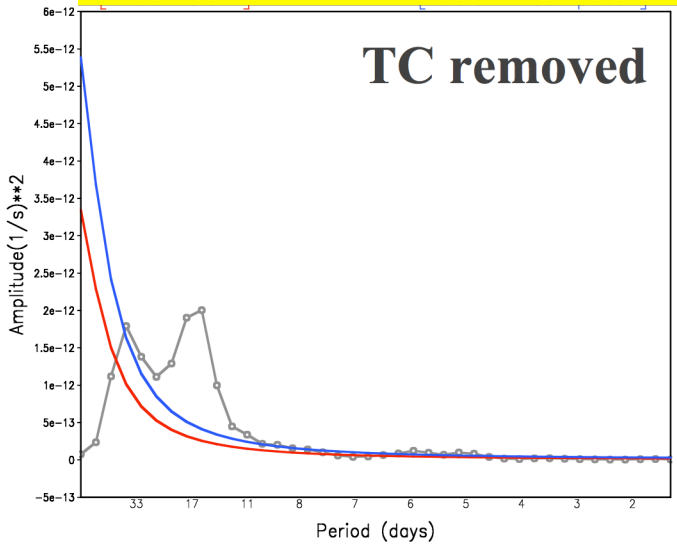
from: An-Kai Lo



1993



TC activity contributes half of intraseasonal variance due to the clustering effect of ISO on TC.



from: An-Kai Lo

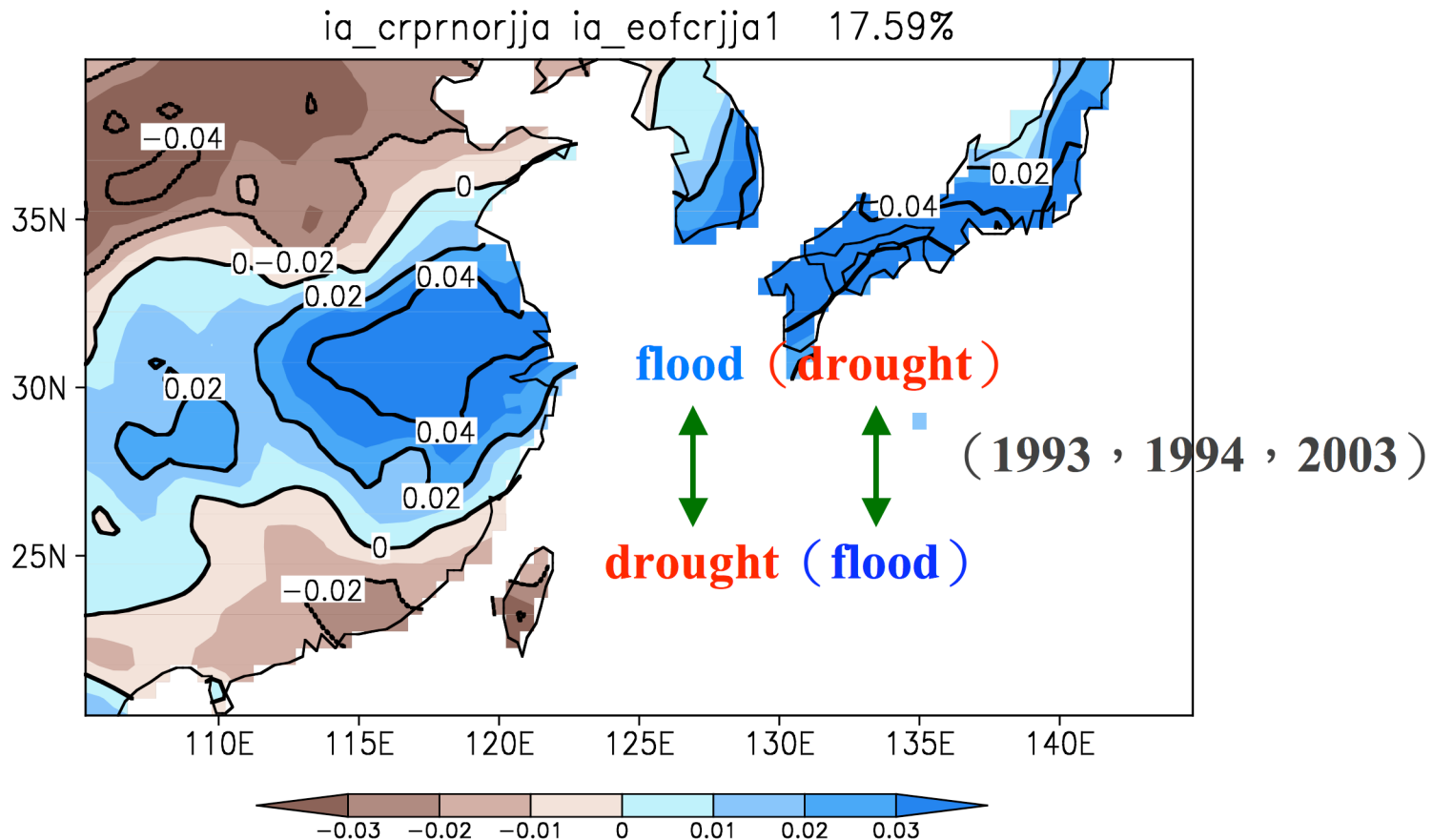
The ‘Tri-pole Rainfall Pattern’

**- Interannual Variability of East Asian Summer
rainfall**

H.-H. Hsu and S.-M. Lin (2005)

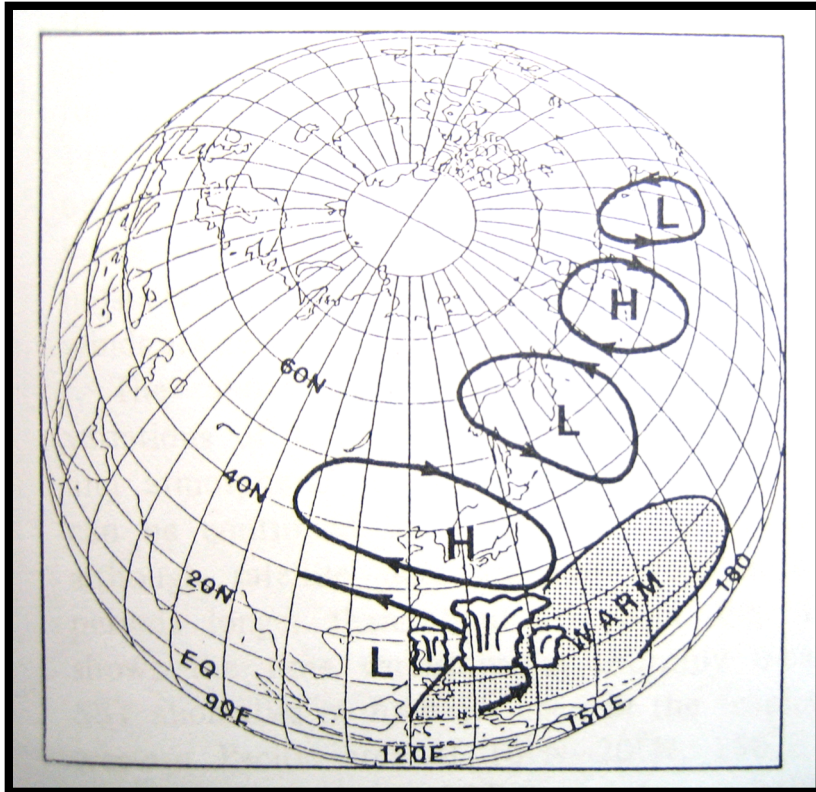
East Asia rainfall EOF1

- Explain 18% of summer rainfall variance
Relatively wet in central China, Korea, and Japan
dry in northern/southern China and Taiwan
- Negative phase: spatial pattern in **opposite** signs



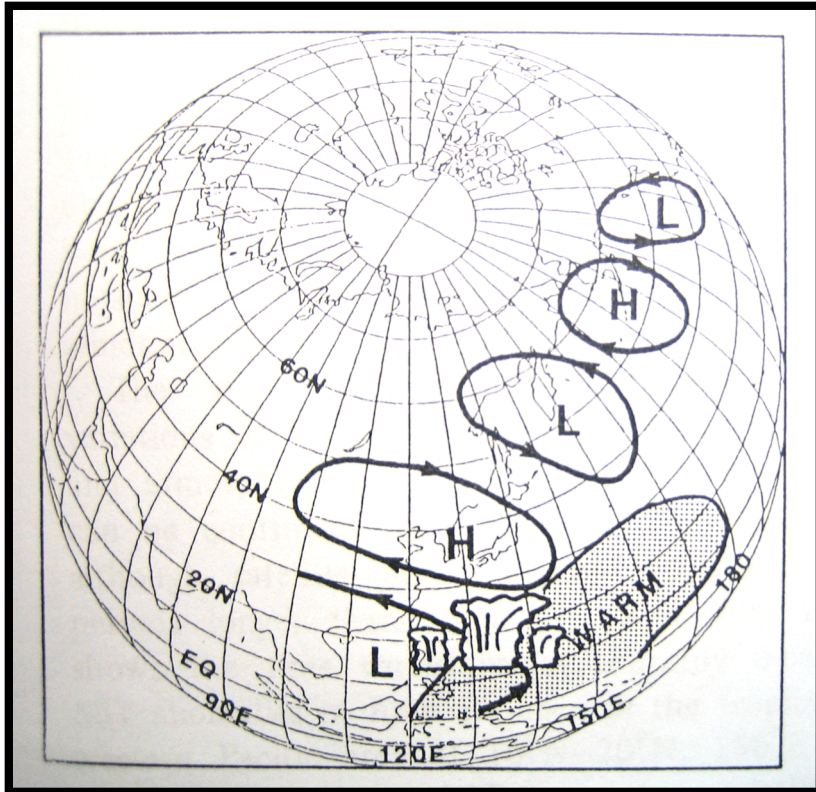
Nitta (1987) ...

Pacific Japan Pattern



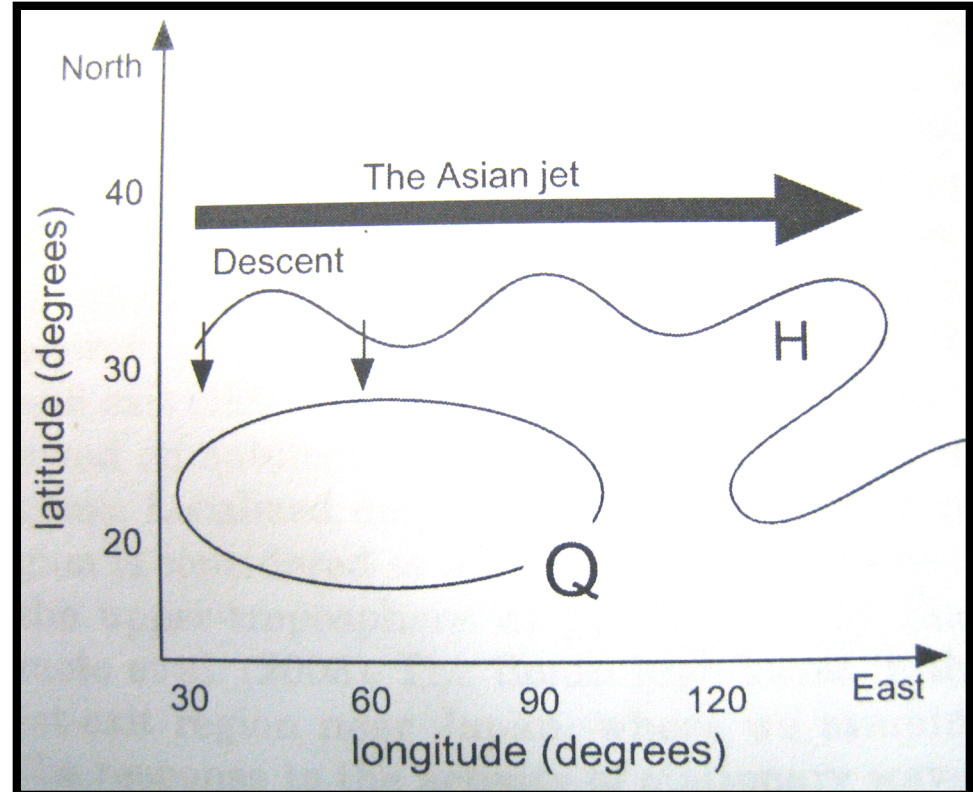
Nitta (1987) ...

Pacific Japan Pattern



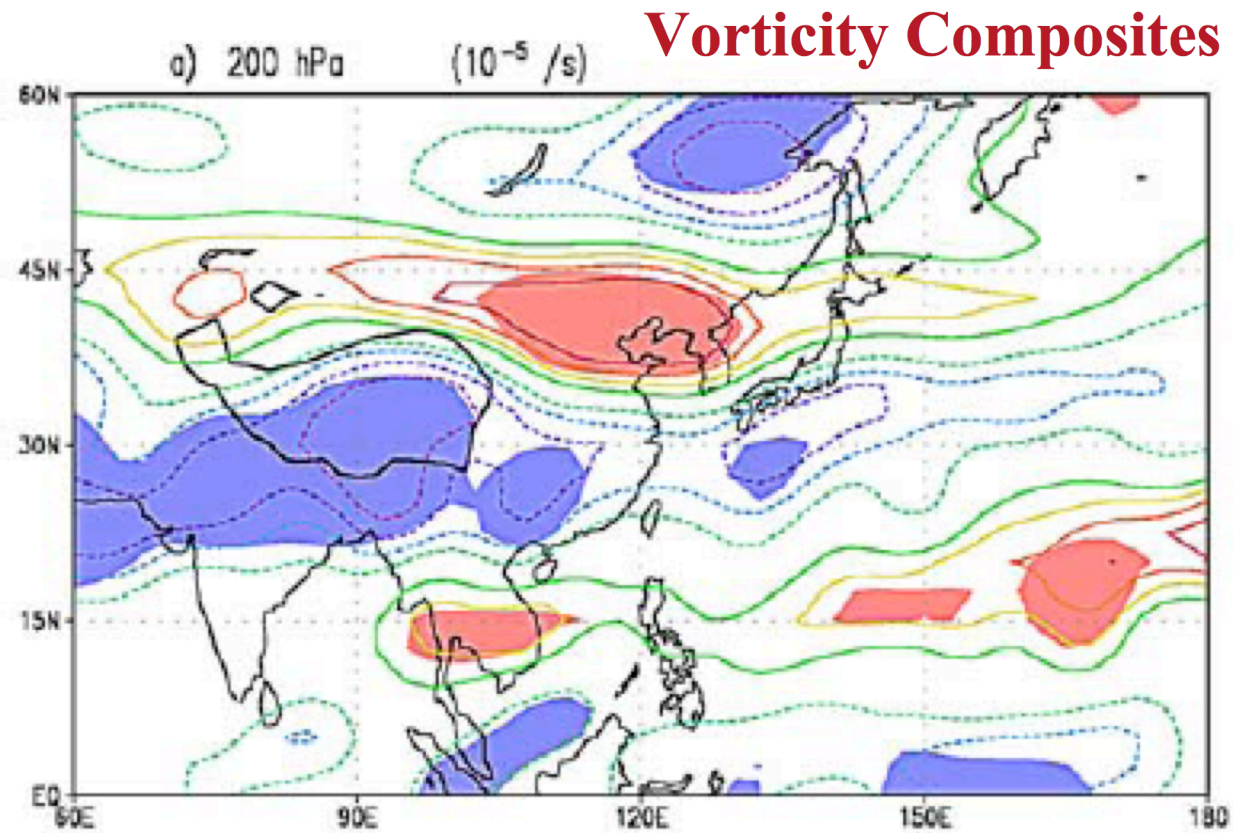
Enomoto et al. (2003) ...

Silk Road Pattern



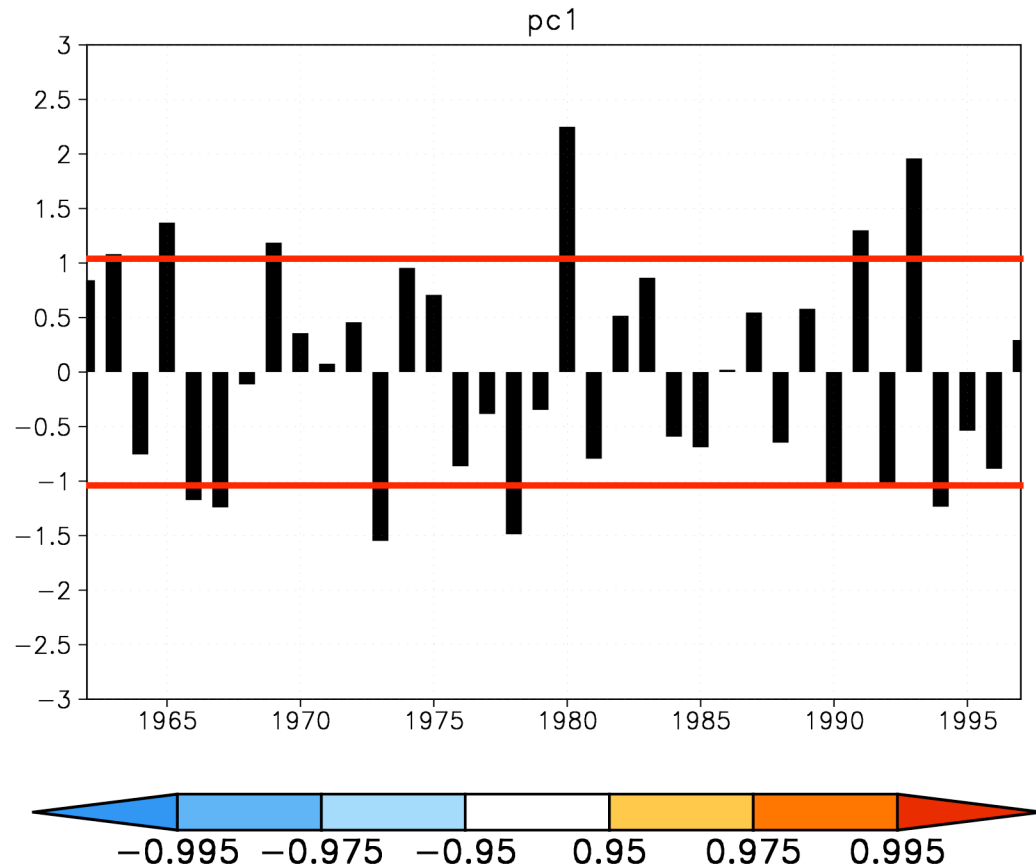
Wave-like perturbations forced by the Tibetan heating?

Hsu and Liu (2003)

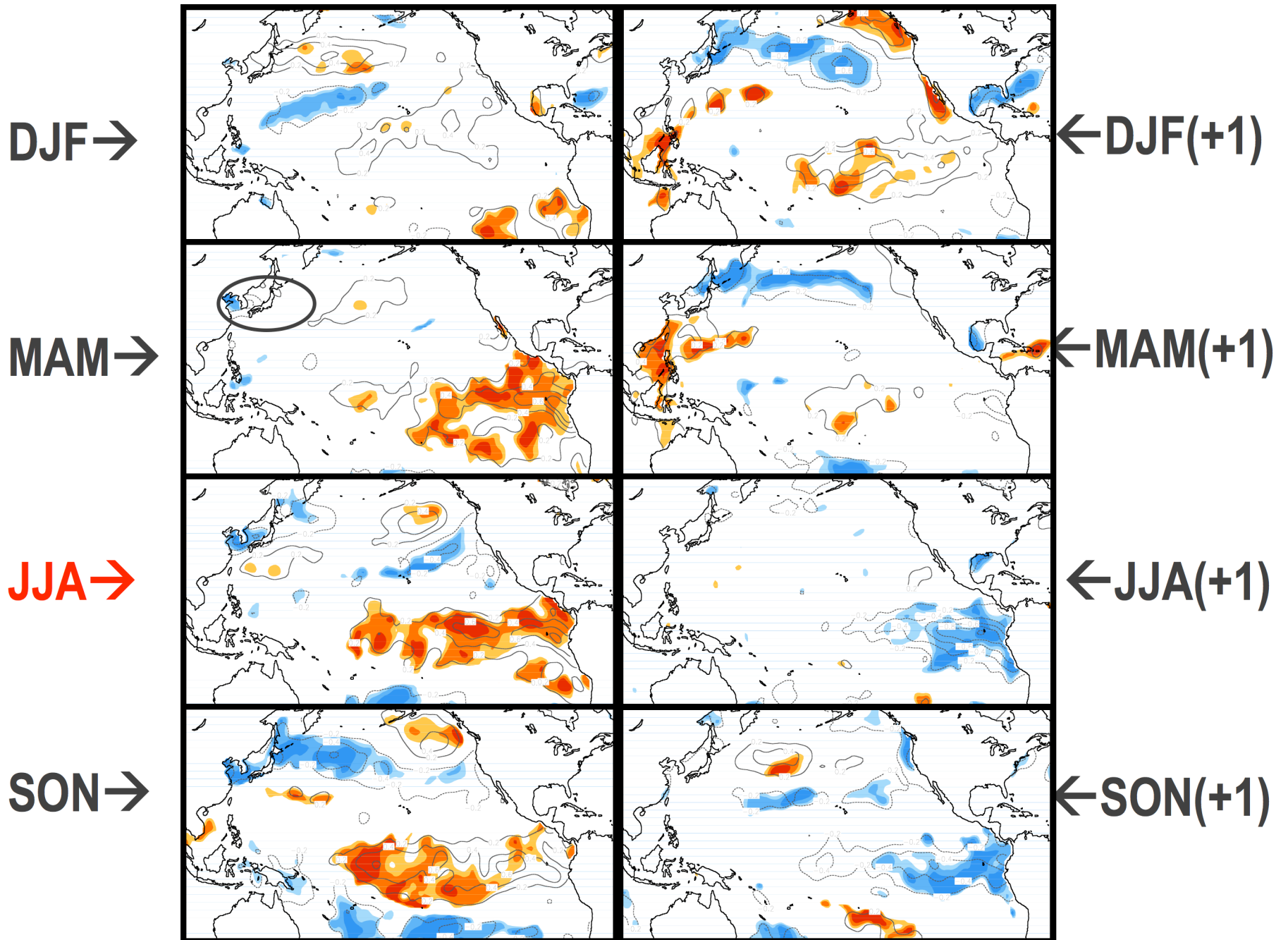


Composite Method

- Positive phase > 1 std \rightarrow wet years (central China)
Negative phase < -1 std \rightarrow dry years
- Averaged anomalies for wet and dry years
 \rightarrow Student's t-test

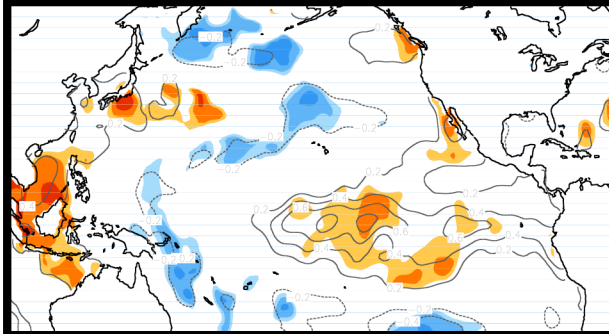


SSTA: **wet** year

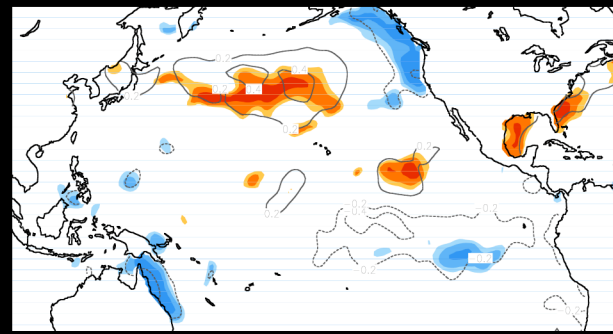


SSTA: **dry** year

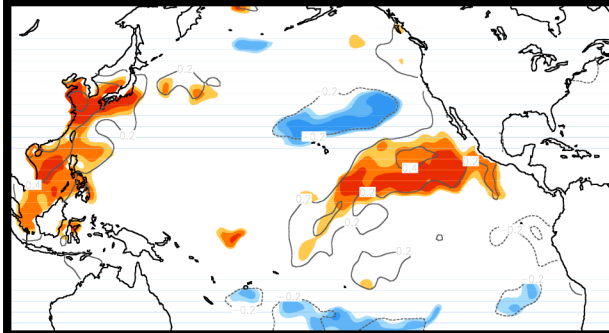
DJF →



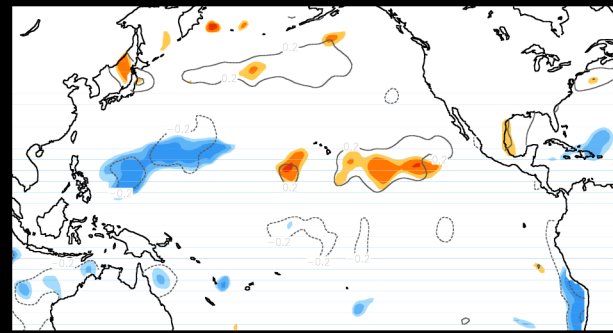
← DJF(+1)



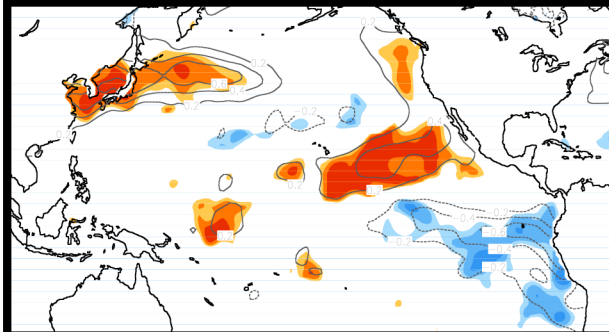
MAM →



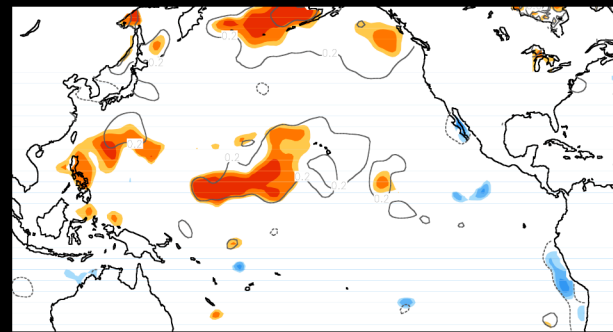
← MAM(+1)



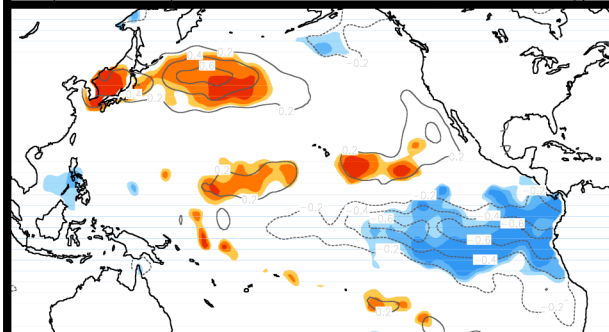
JJA →



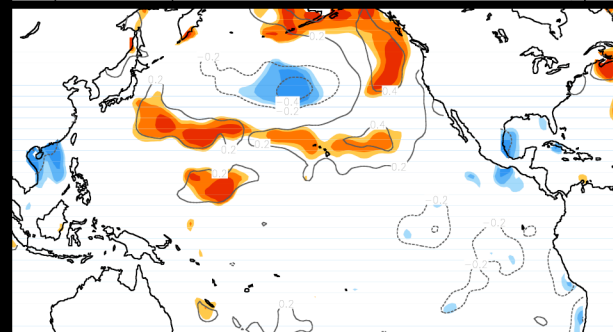
← JJA(+1)



SON →

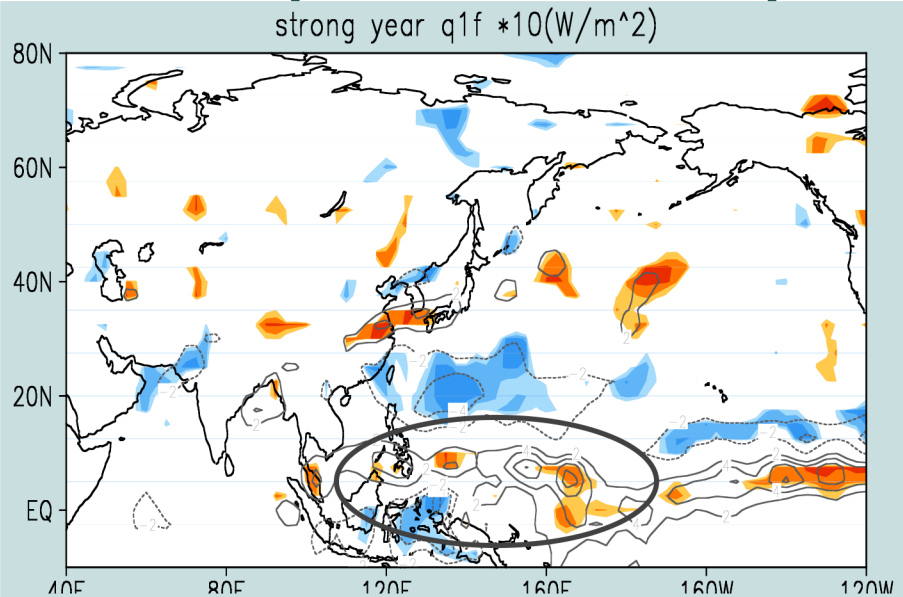


← SON(+1)

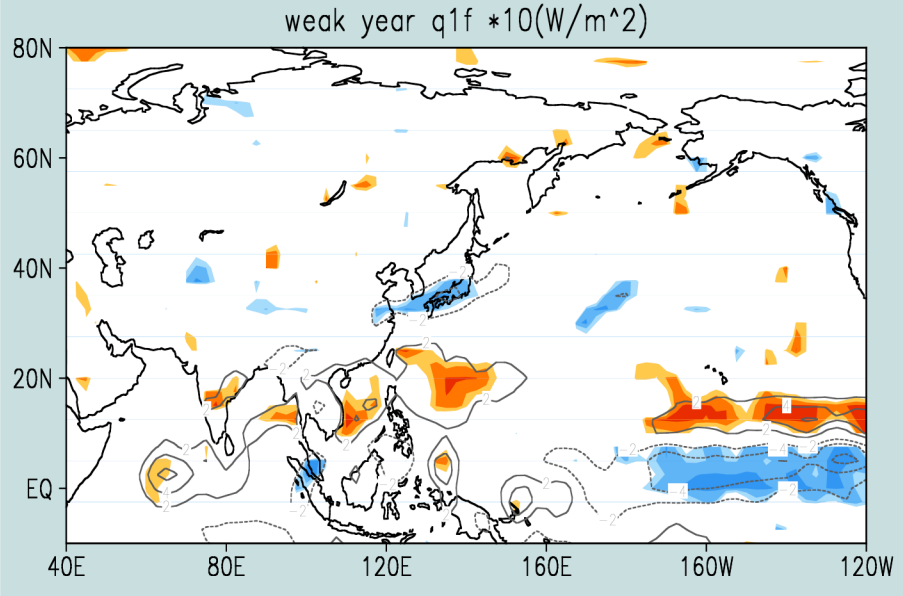


Q1 (Heat Source)

WET



DRY



Summary

- Wet-year SSTA exhibits El Nino (year 0-1) characteristics, while dry-year SSTA exhibits only weak La Nina (year 0-1) characteristics.
- Relationship between the SST and diabatic heating anomalies:
 - Tropical Eastern Pacific:
 - + (-) SSTA \leftrightarrow + (-) heating anomalies
 - SSTA affects atmosphere (heating anomalies)
 - Extratropical W. Pacific:
 - + (-) SSTA \leftrightarrow - (+) heating anomalies
 - Atmosphere affects SSTA

Vorticity*

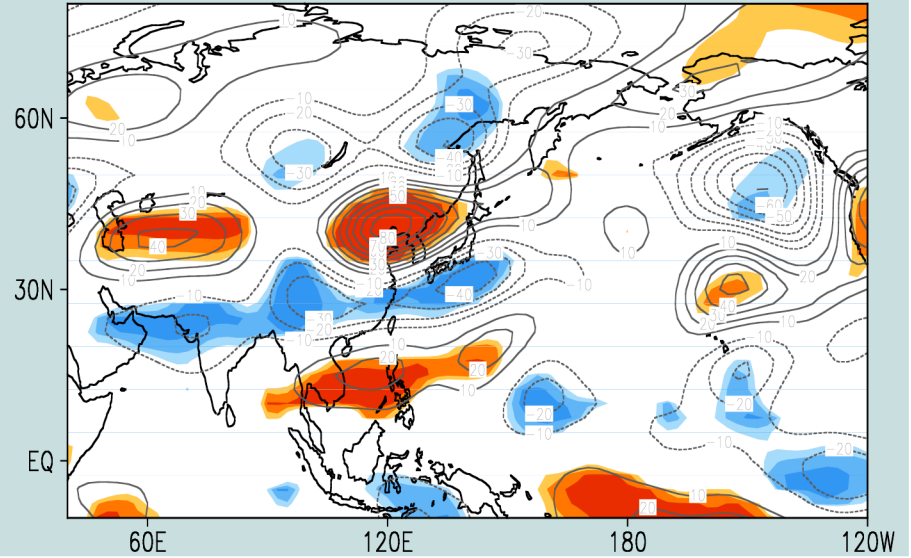
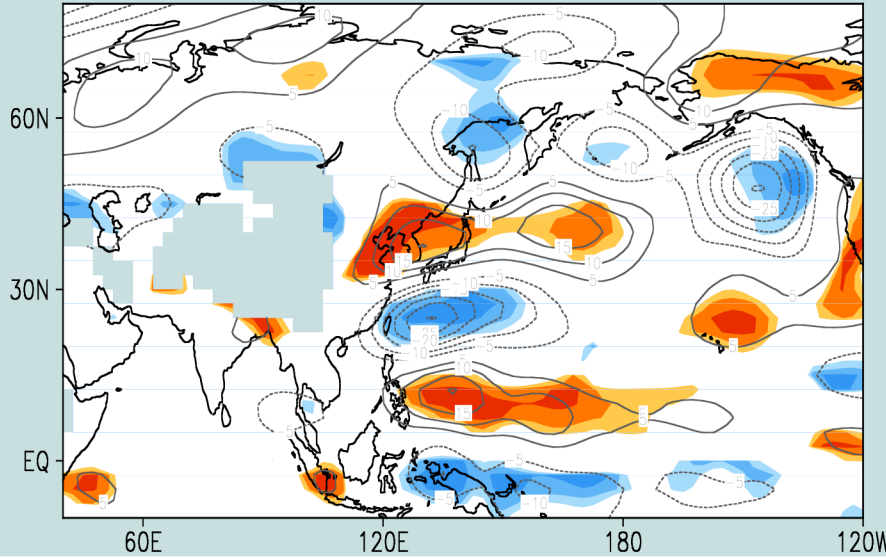
850 hPa

200 hPa

strong year vor $\cdot 10^{-7}(1/s)$ 850hPa

WET

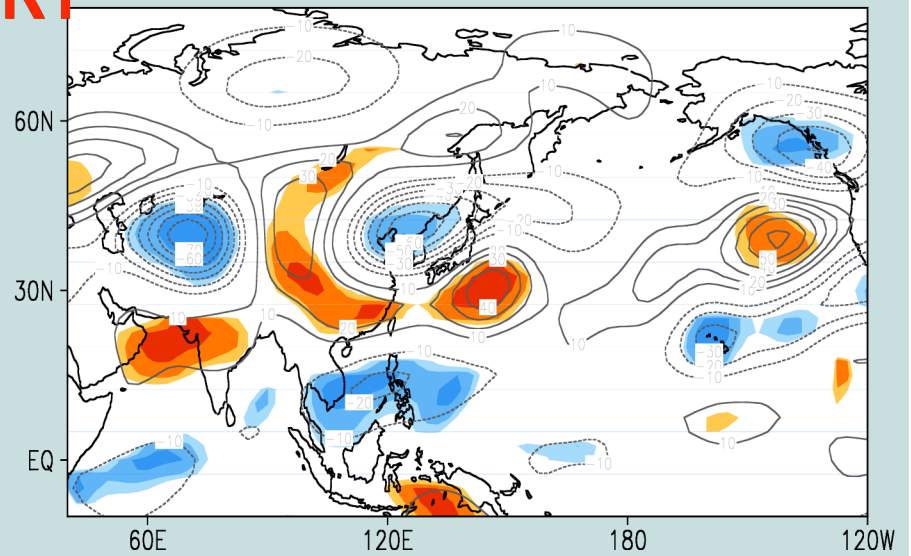
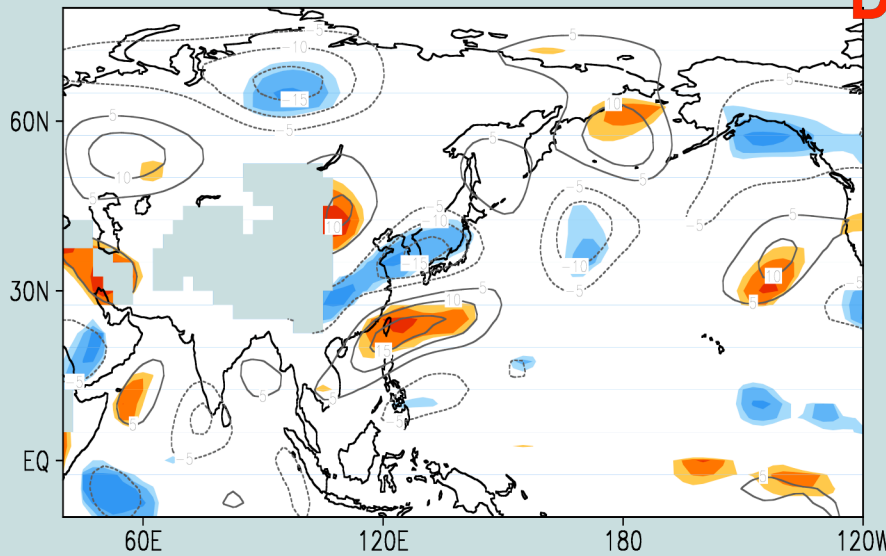
strong year vor $\cdot 10^{-7}(1/s)$ 200hPa



weak year vor $\cdot 10^{-7}(1/s)$ 850hPa

DRY

weak year vor $\cdot 10^{-7}(1/s)$ 200hPa



Vorticity*

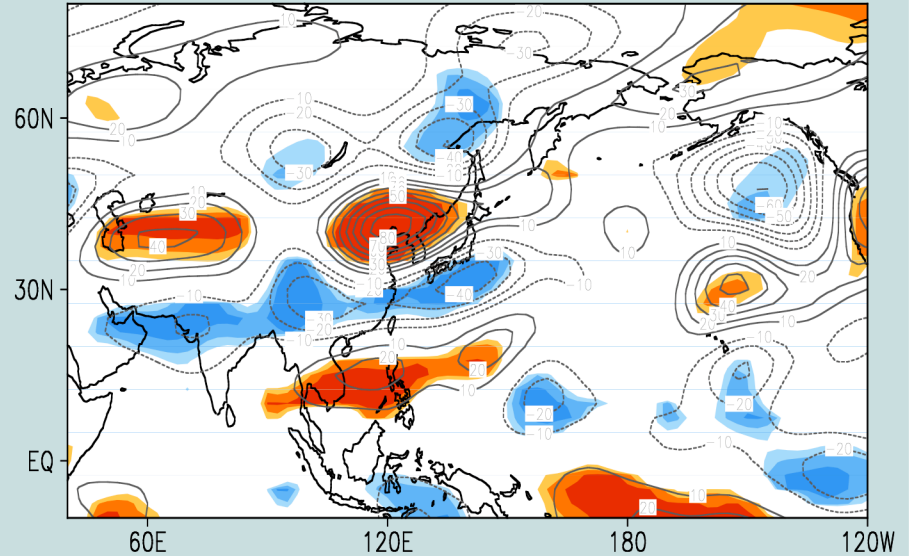
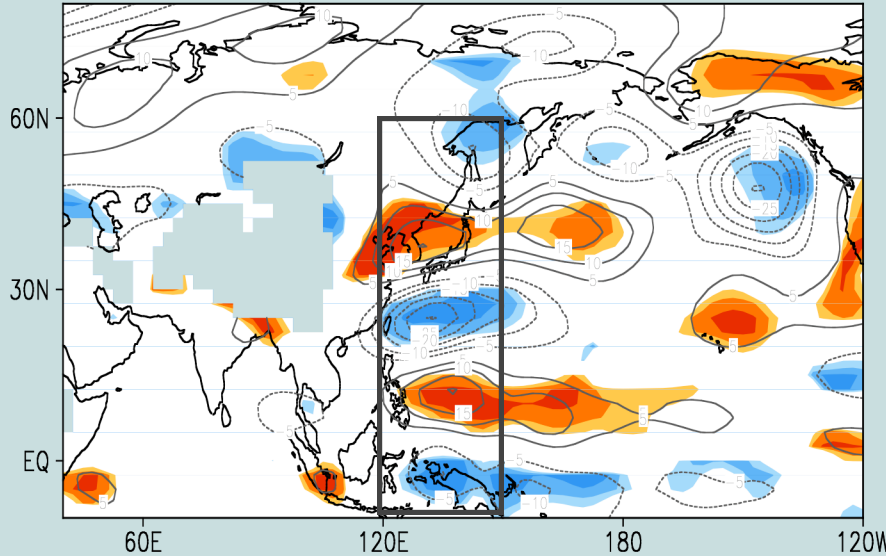
850 hPa

200 hPa

strong year vor $\times 10^{-7}(1/s)$ 850hPa

WET

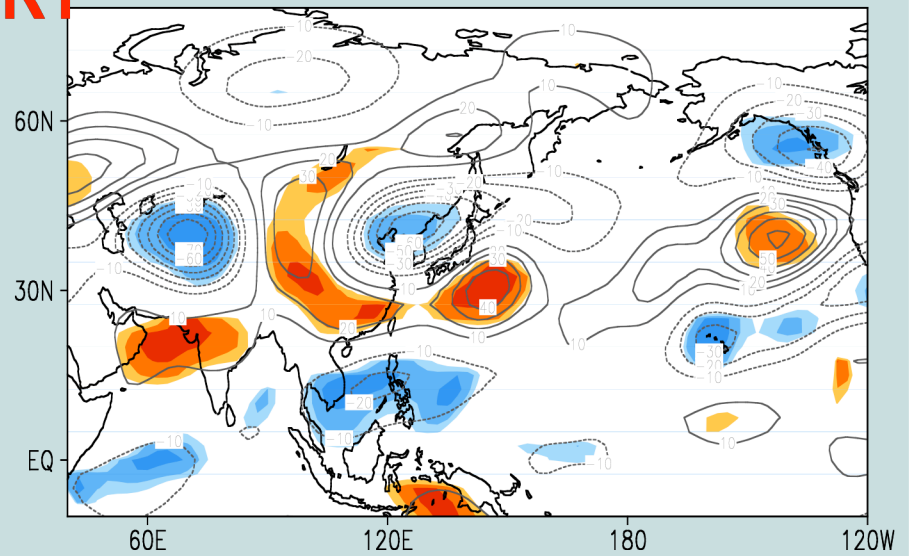
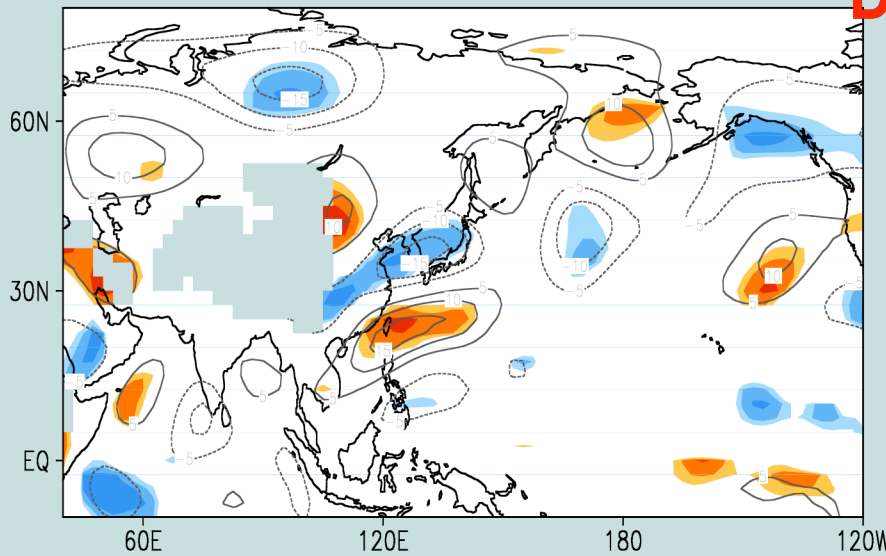
strong year vor $\times 10^{-7}(1/s)$ 200hPa



weak year vor $\times 10^{-7}(1/s)$ 850hPa

DRY

weak year vor $\times 10^{-7}(1/s)$ 200hPa



Vorticity*

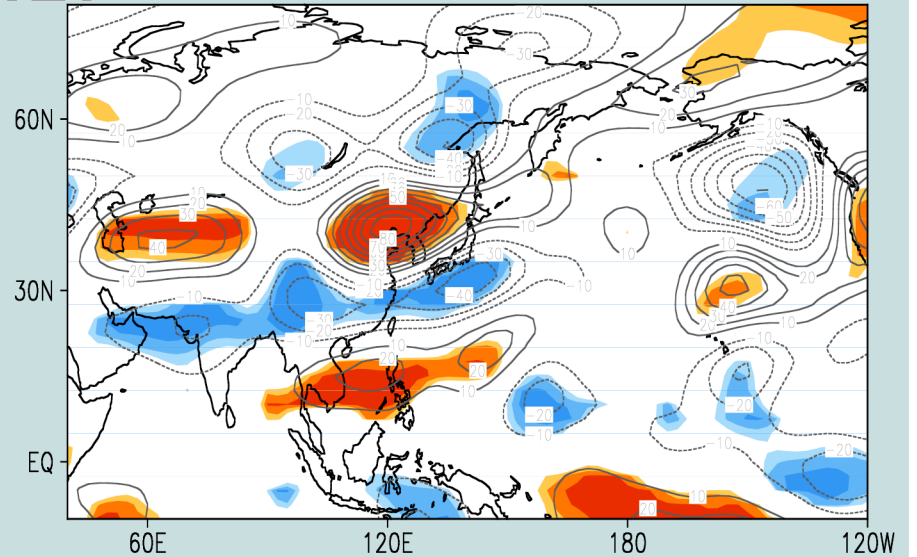
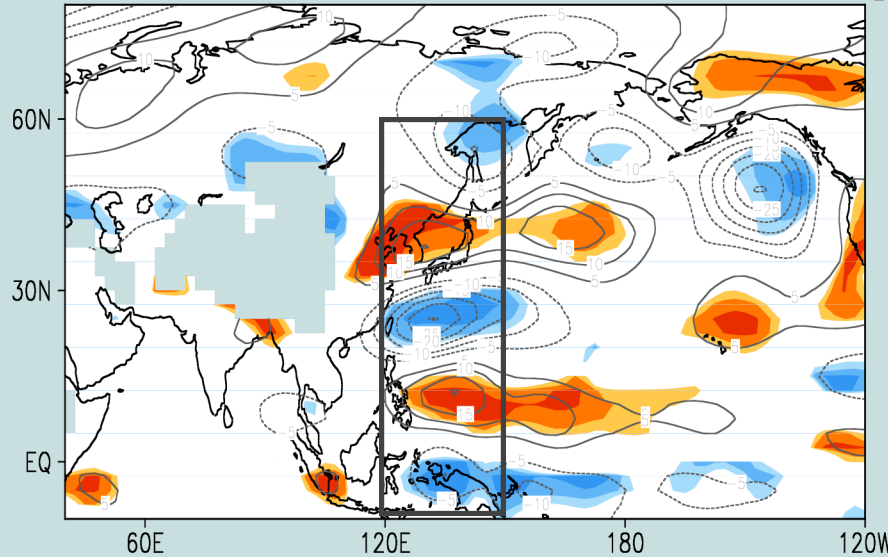
850 hPa

200 hPa

WET

strong year vor $\times 10^{-7}(1/s)$ 850hPa

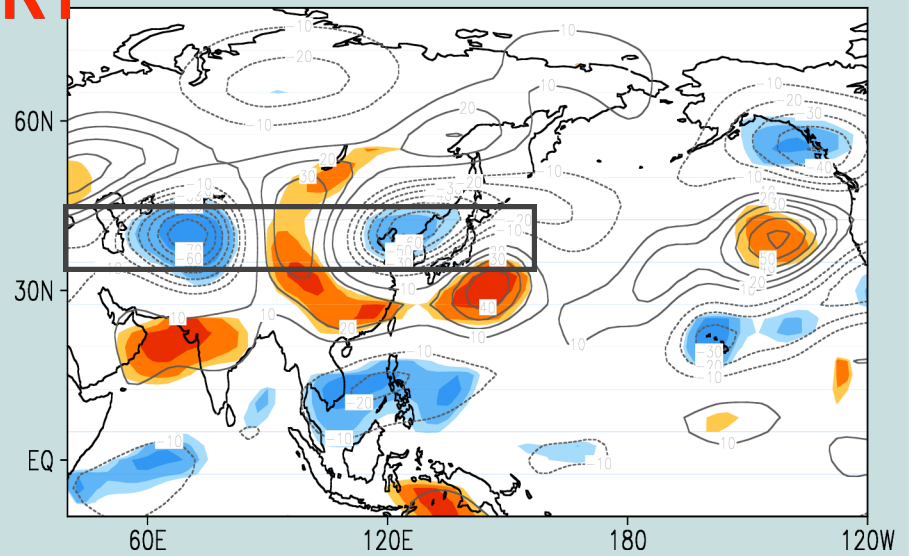
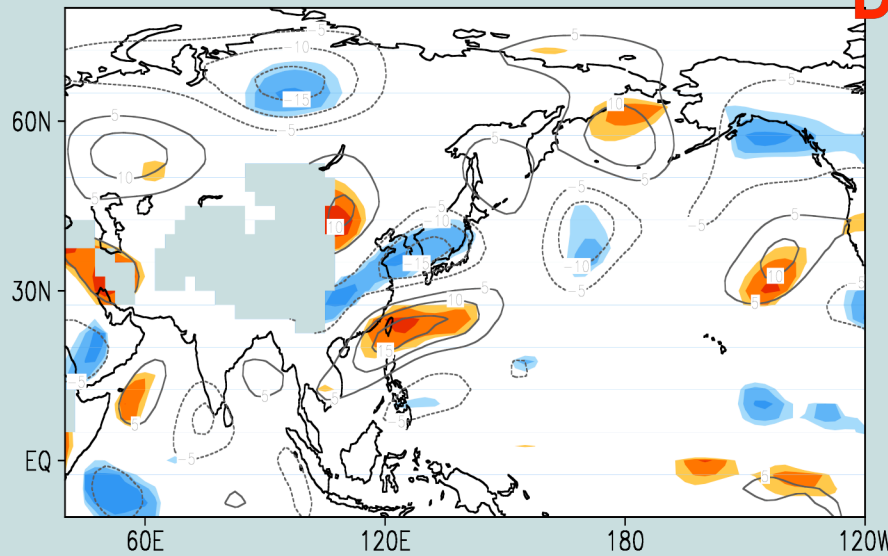
strong year vor $\times 10^{-7}(1/s)$ 200hPa



DRY

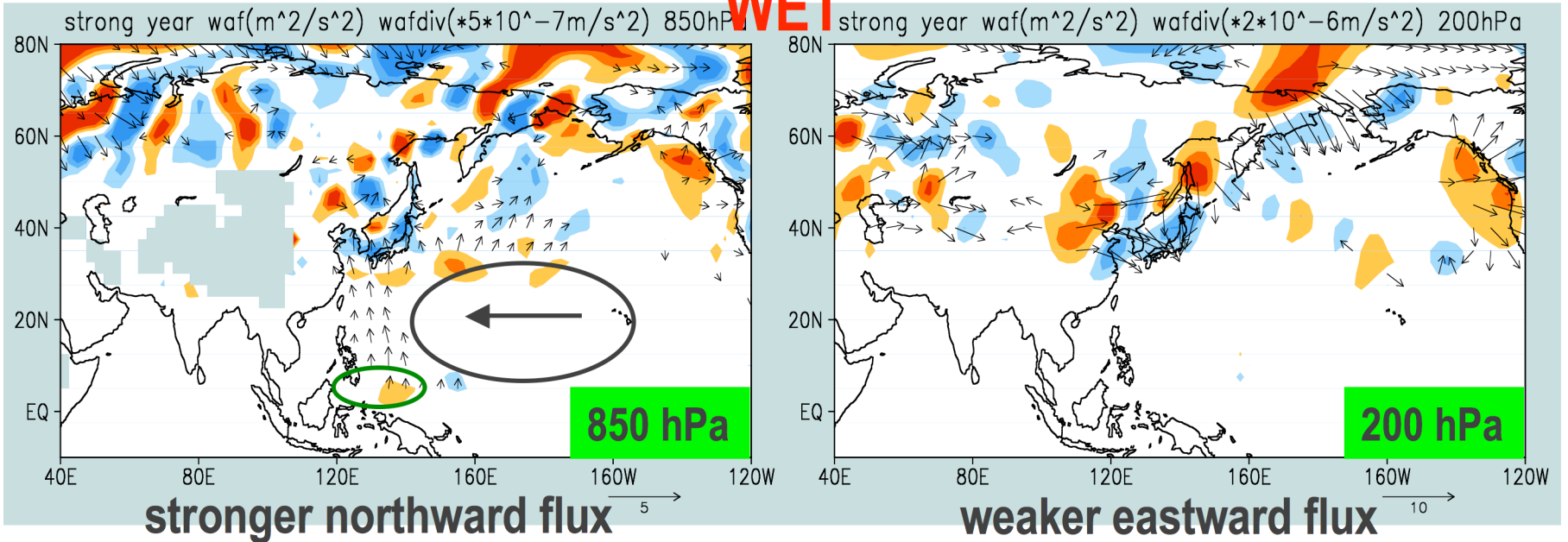
weak year vor $\times 10^{-7}(1/s)$ 850hPa

weak year vor $\times 10^{-7}(1/s)$ 200hPa

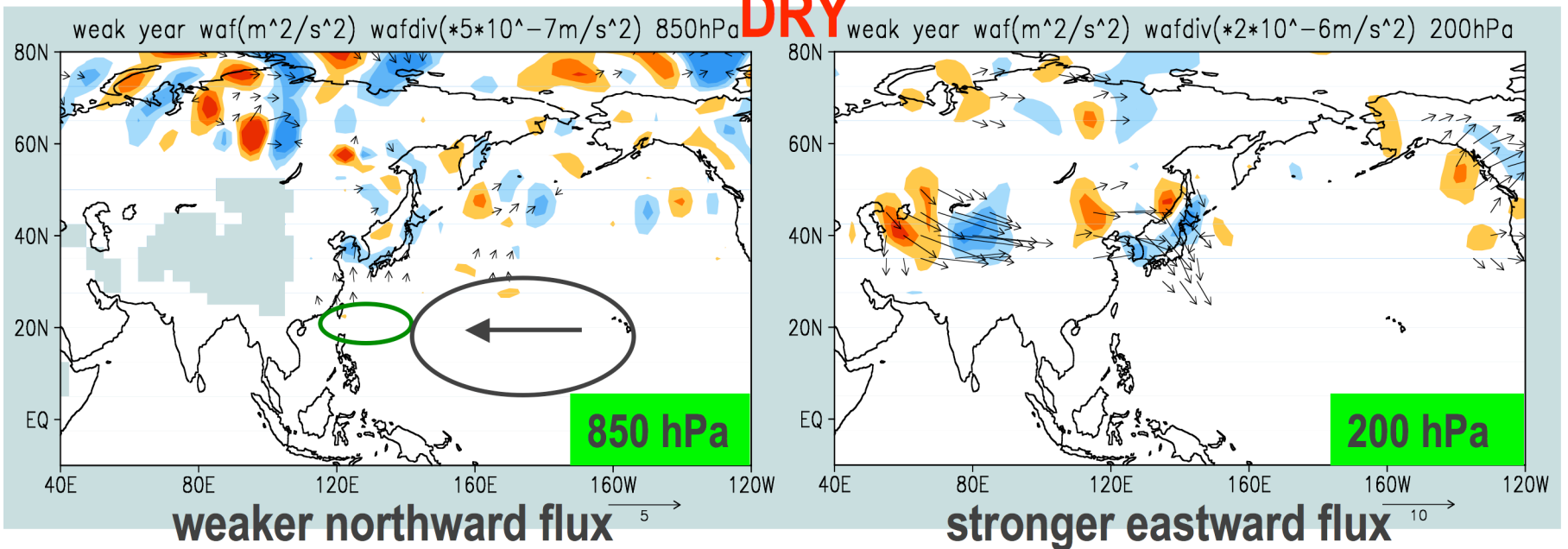


Wave Activity Flux (Takaya and Nakamura 1997)

WET



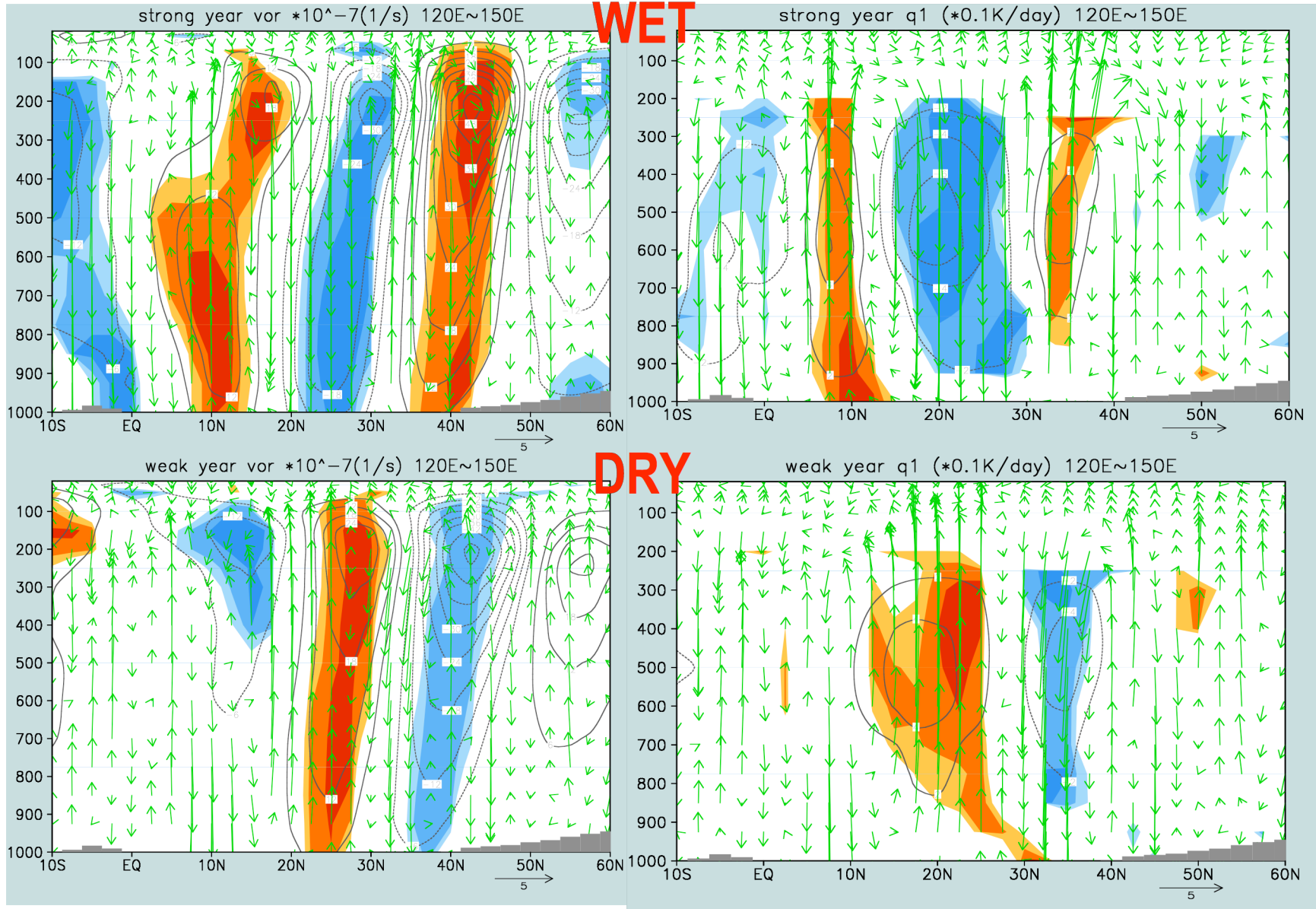
DRY



vorticity

120E-150E

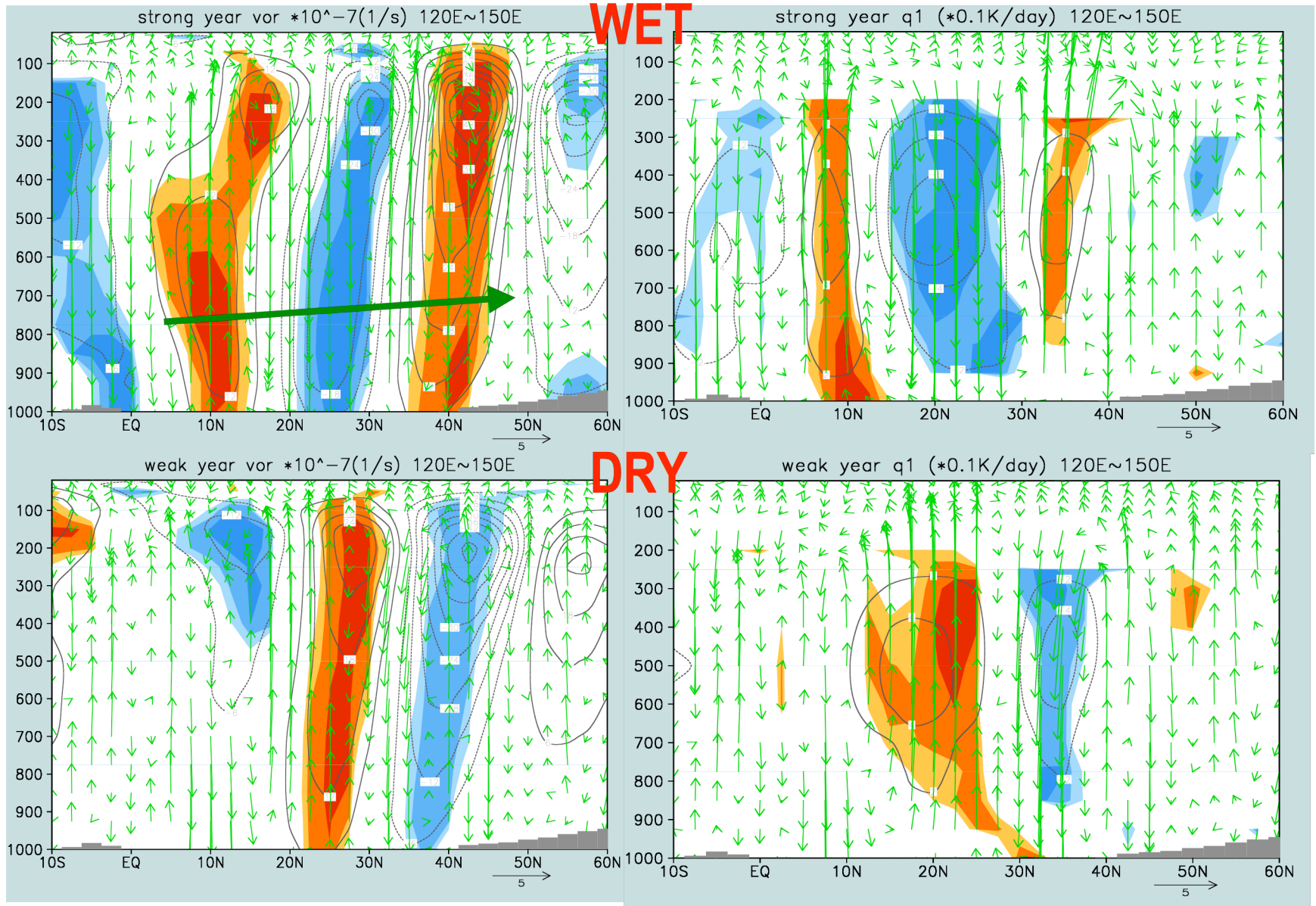
Q1



vorticity

120E-150E

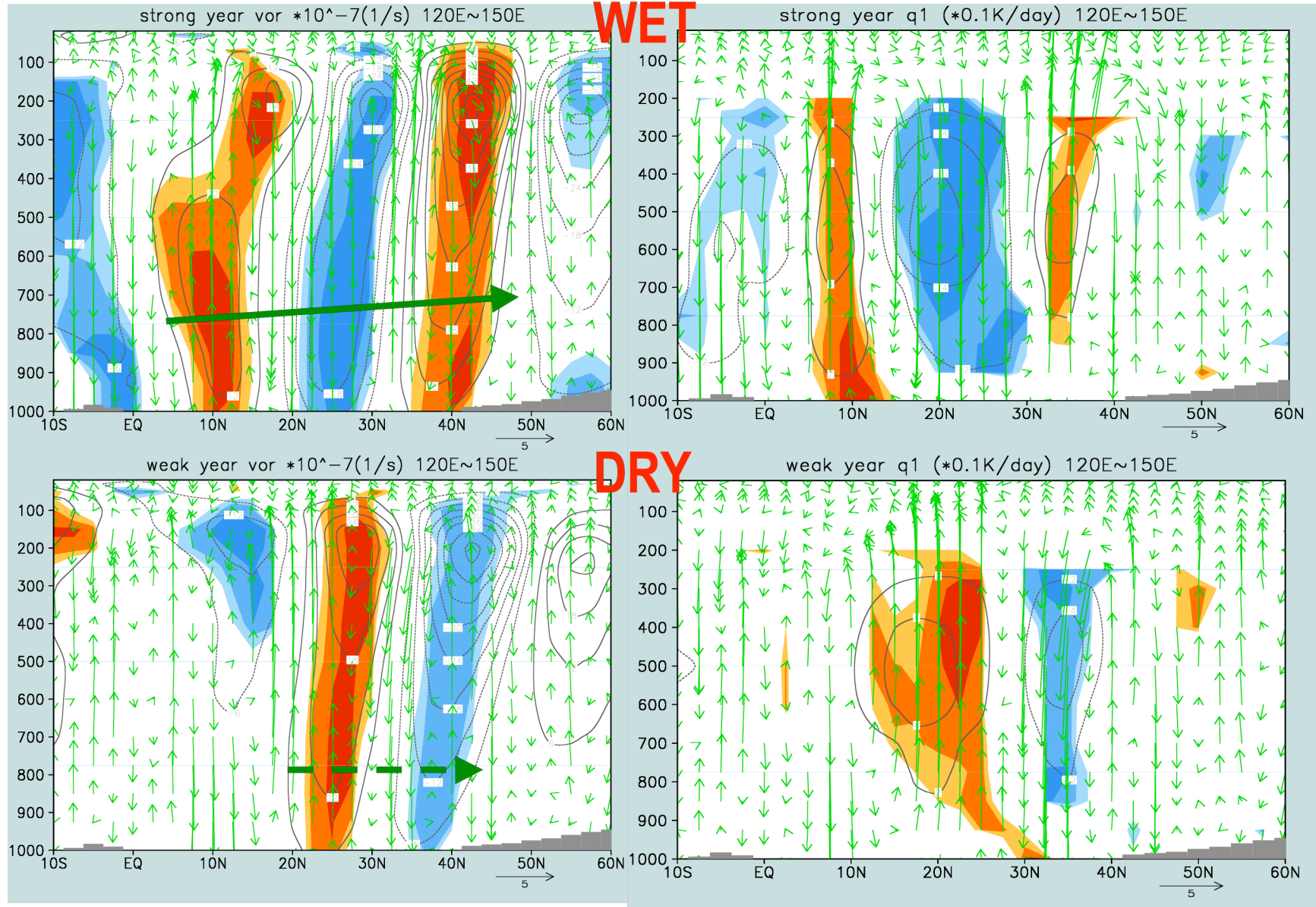
Q1



vorticity

120E-150E

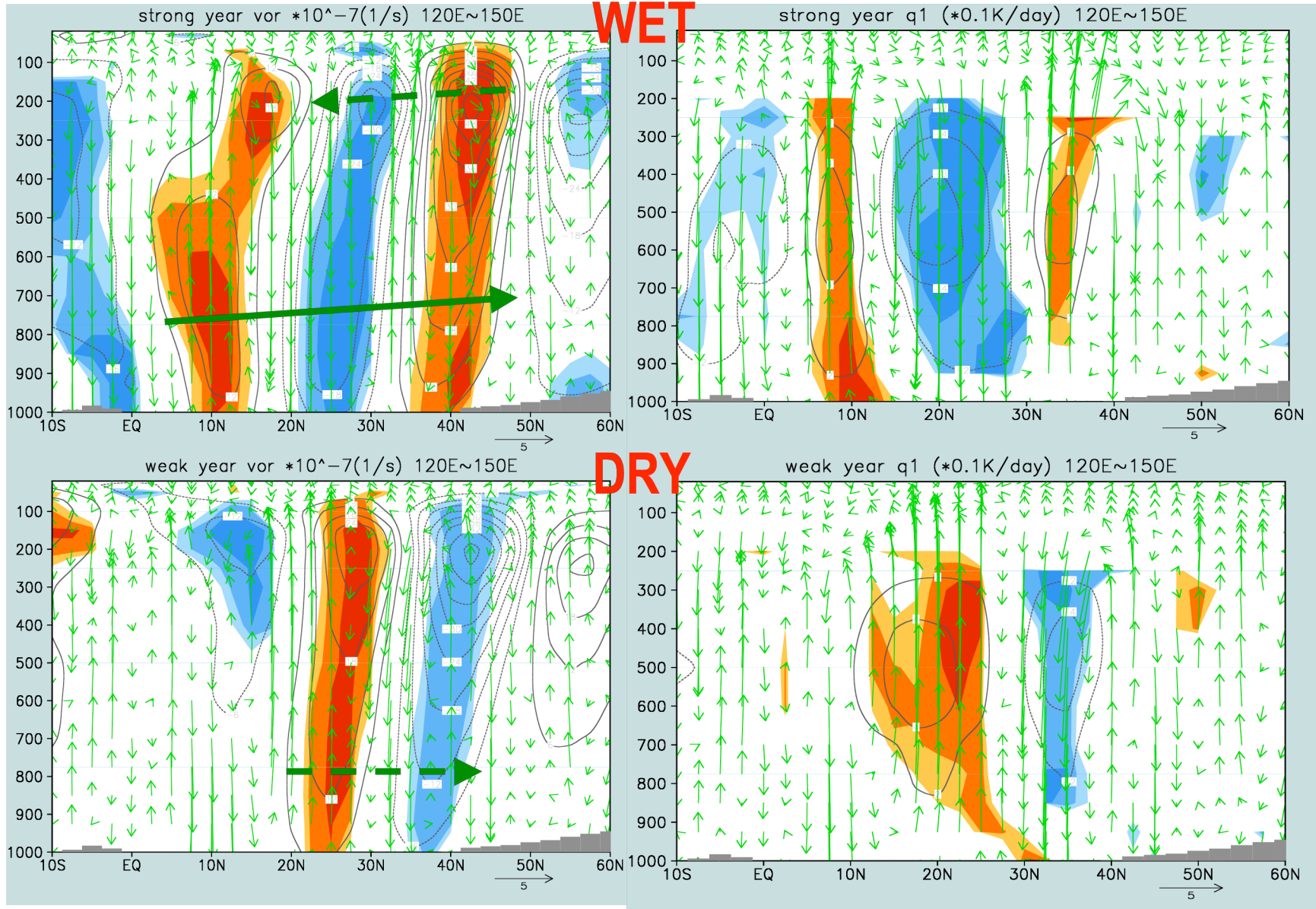
Q1



vorticity

120E-150E

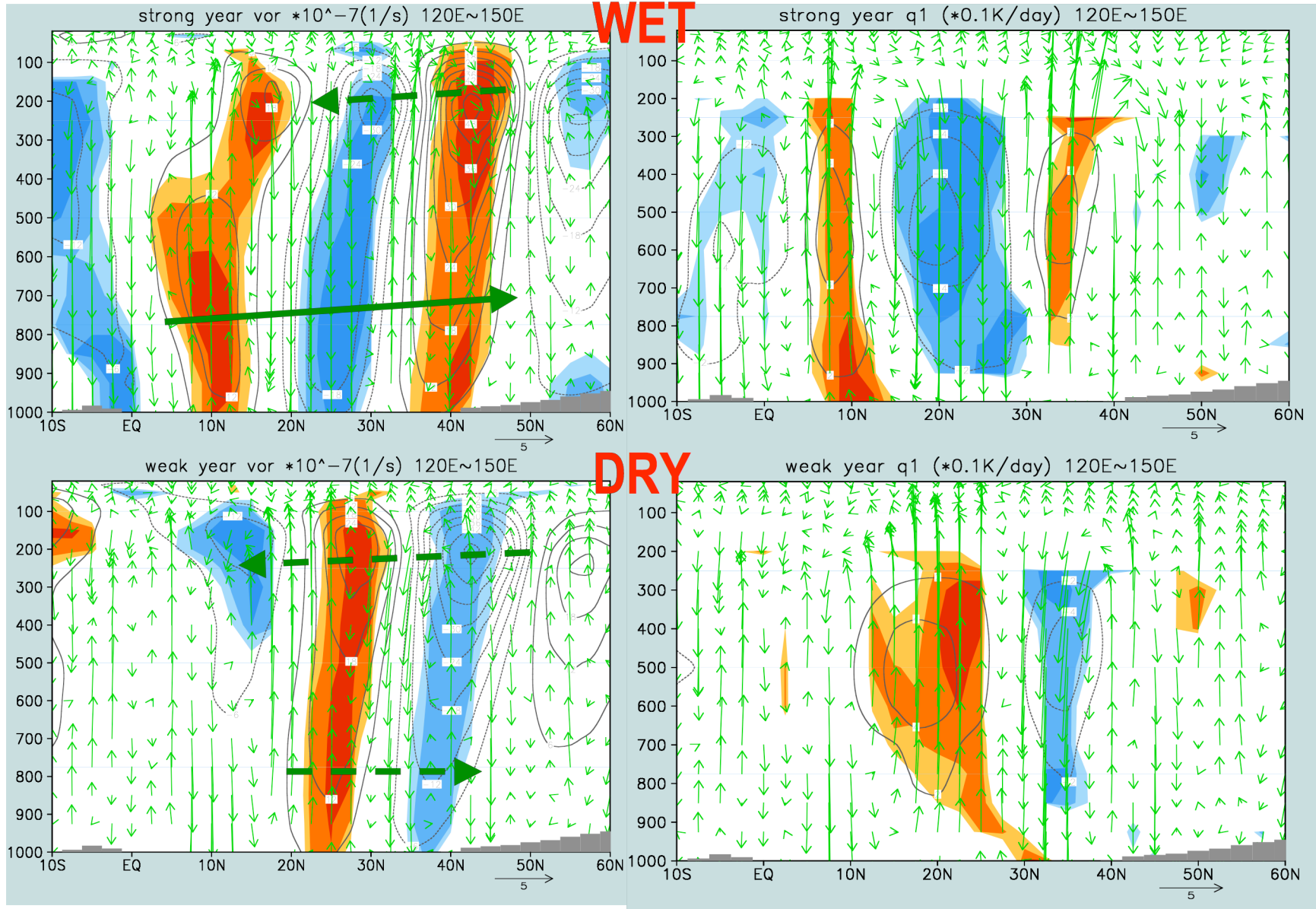
Q1



vorticity

120E-150E

Q1

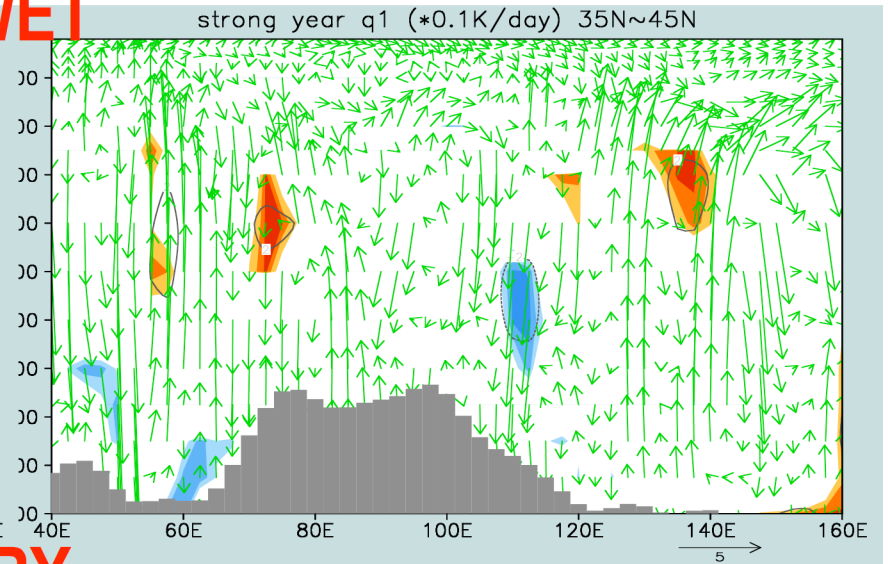
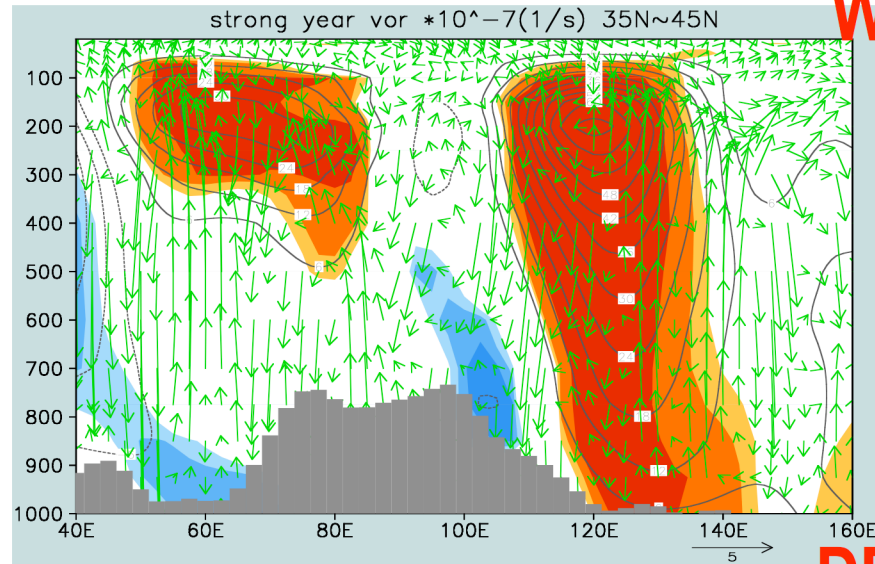


vorticity

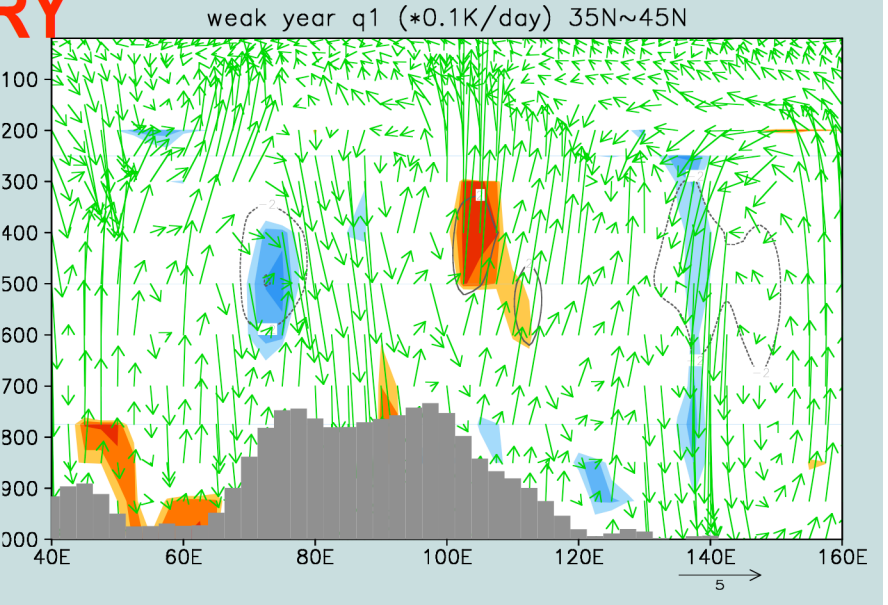
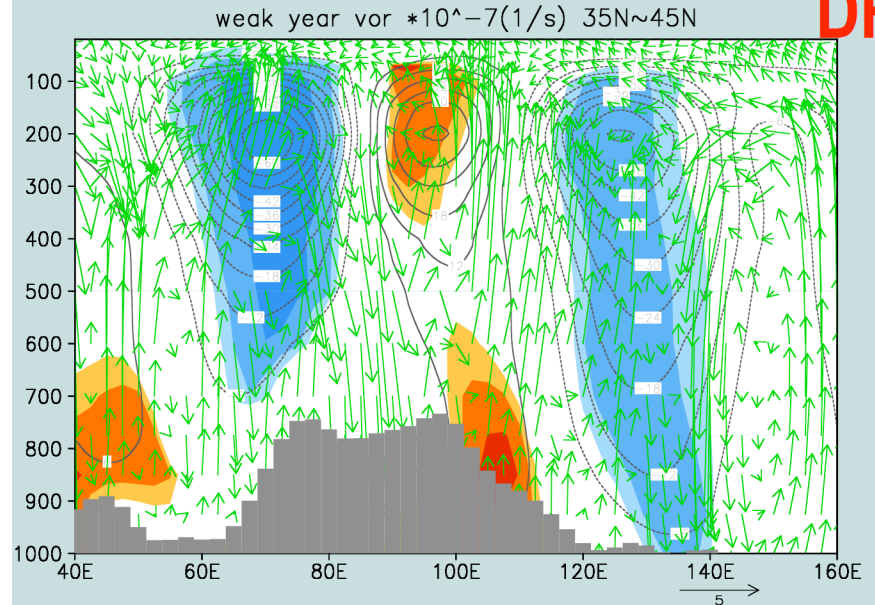
35N-45N

Q1

WET



DRY

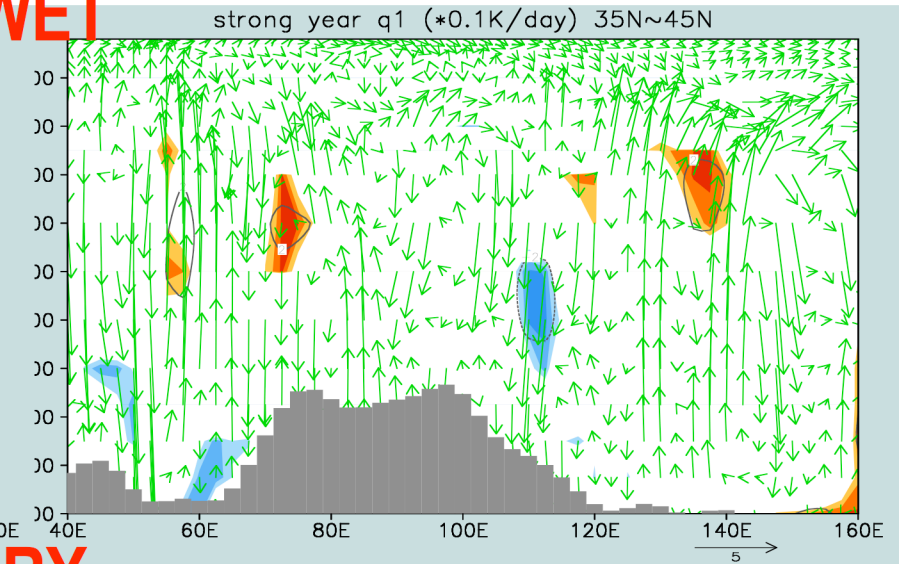
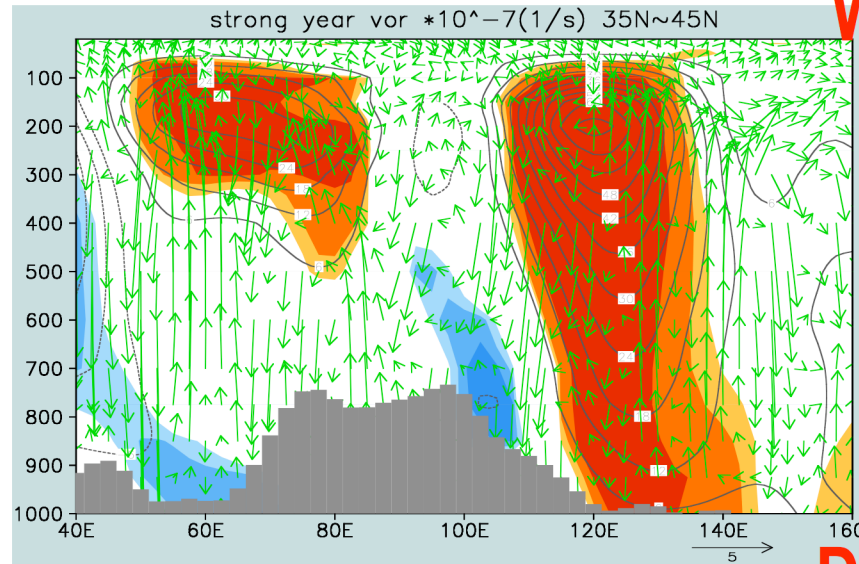


vorticity

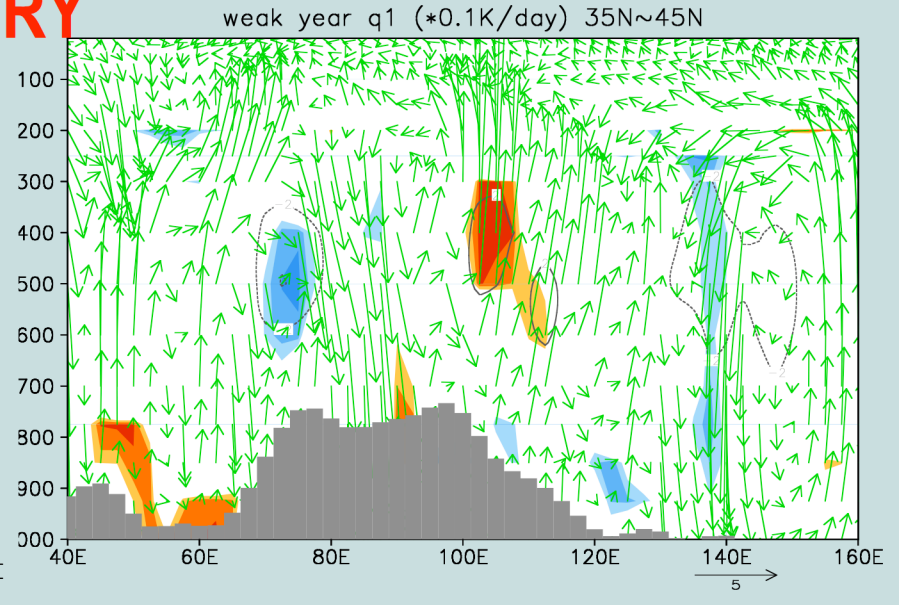
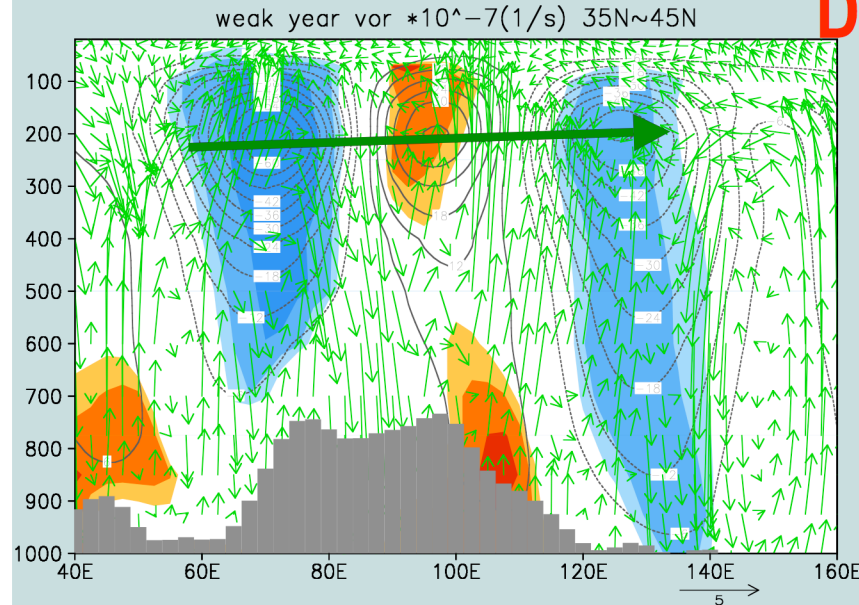
35N-45N

Q1

WET



DRY

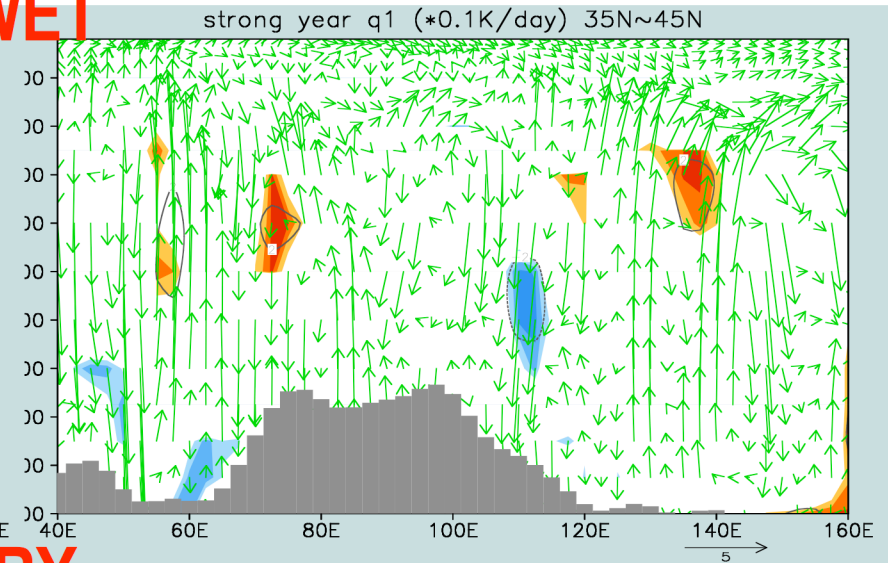
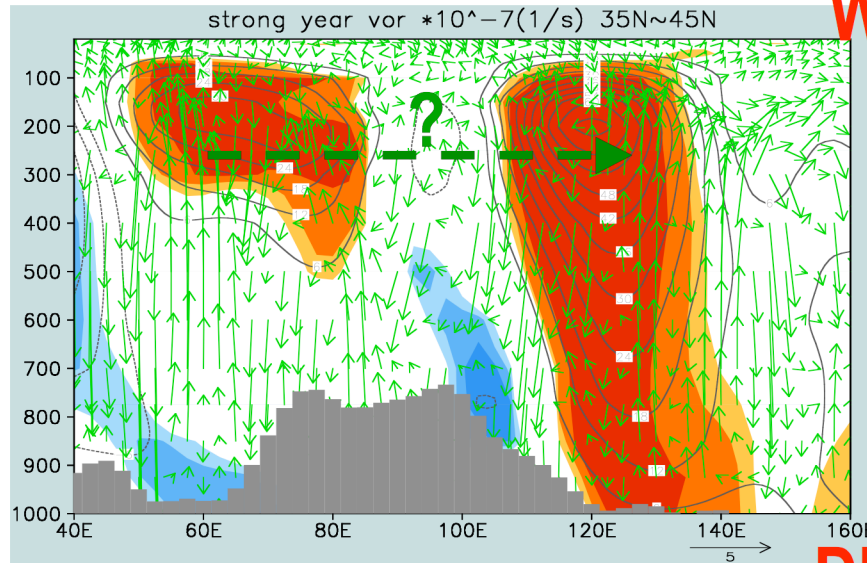


vorticity

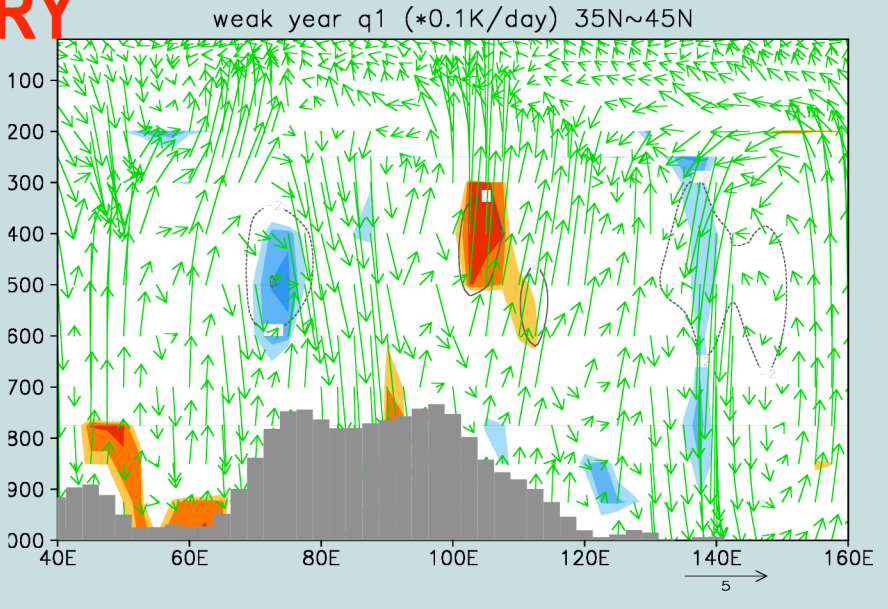
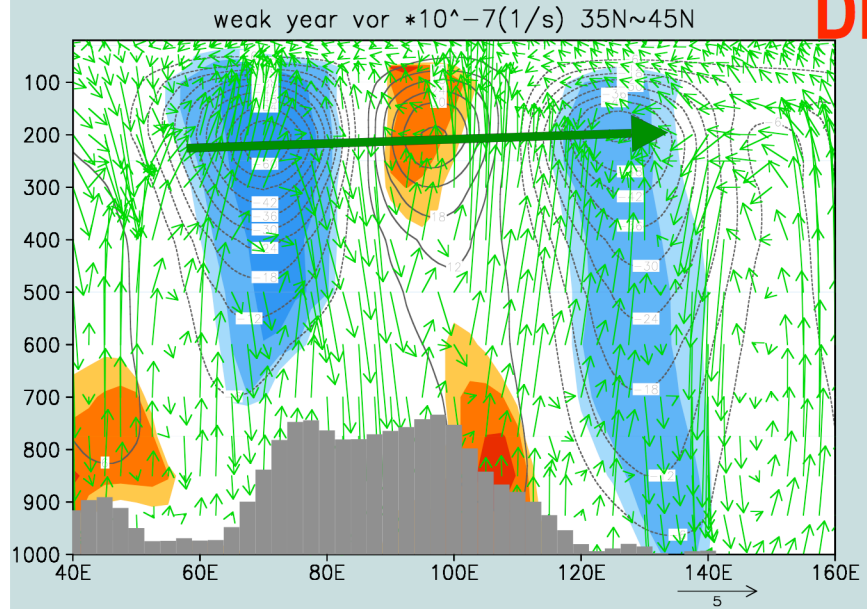
35N-45N

Q1

WET



DRY



Summary

- Two dominant wave-like patterns:
 1. North-south wave train along the E. Asian coast.
 2. East-west wave train over the extratropical Eurasian continent in the 30N-50N latitudinal band.
- Both wave trains exhibit slightly tilting vertical structure.

Summary (cont.)

North-South wave train

- Similar to the Pacific Japan pattern, but not necessarily forced by tropical heating
- Energy propagation: **northward** in the lower troposphere and **southward** tendency in the upper troposphere
- Northward propagation forced by tropical heating is **confined** in a narrow westerly zone along the East Asian coast.
- Tropical **heating** in the tropical western Pacific is an important factor affecting **rainfall in central China and Japan in wet years**.

Summary (cont.)

East-west extratropical wave train

- Similar to Silk Road pattern
- Eastward energy propagation
- Stronger in **dry** years

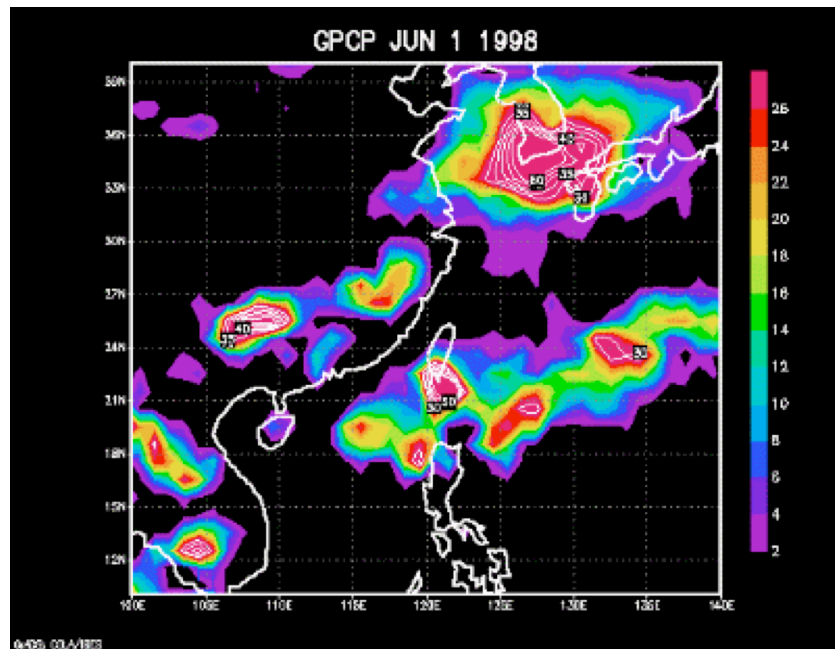
Can not be explained by a simple mechanism:

- Nitta, Enomoto, Hsu and Liu revealed only one part of the whole picture.

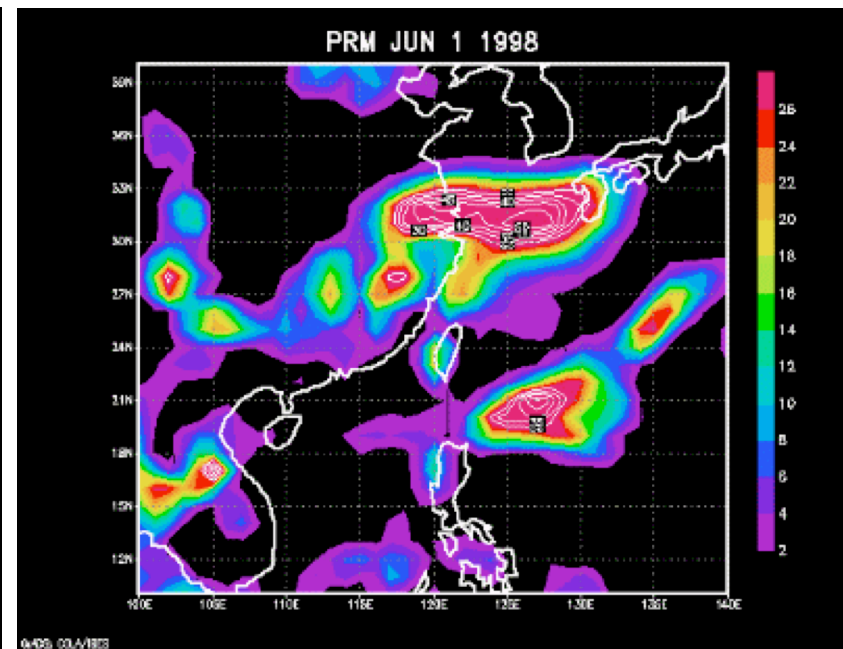
Regional Climate Simulation using Purdue Regional Model

June 1998

observed precip.



simulated precip.



Hsu H.-H., W.-S. Kau, W.-R. Hsu, W.-Y. Sun, Y.-C. Yu, Y.-S. Tong, C.-F. Shi, W.-N. Yu

10 summer simulation

CMAP 1991

CMAP 1992

CMAP 1993

CMAP 1994

CMAP 1995

0.83

0.72

0.64

0.68

0.73

PRM 1991

PRM 1992

PRM 1993

PRM 1994

PRM 1995

CMAP 1996

CMAP 1997

CMAP 1998

CMAP 1999

CMAP 2000

0.81

0.74

0.66

0.88

0.80

PRM 1996

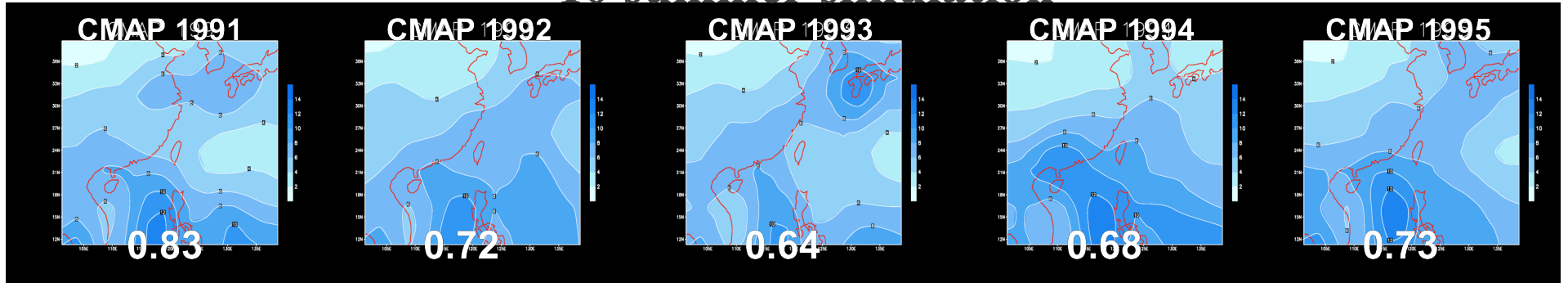
PRM 1997

PRM 1998

PRM 1999

PRM 2000

10 summer simulation



PRM 1991

PRM 1992

PRM 1993

PRM 1994

PRM 1995

CMAP 1996

CMAP 1997

CMAP 1998

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CMAP 2000

0.81

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PRM 1996

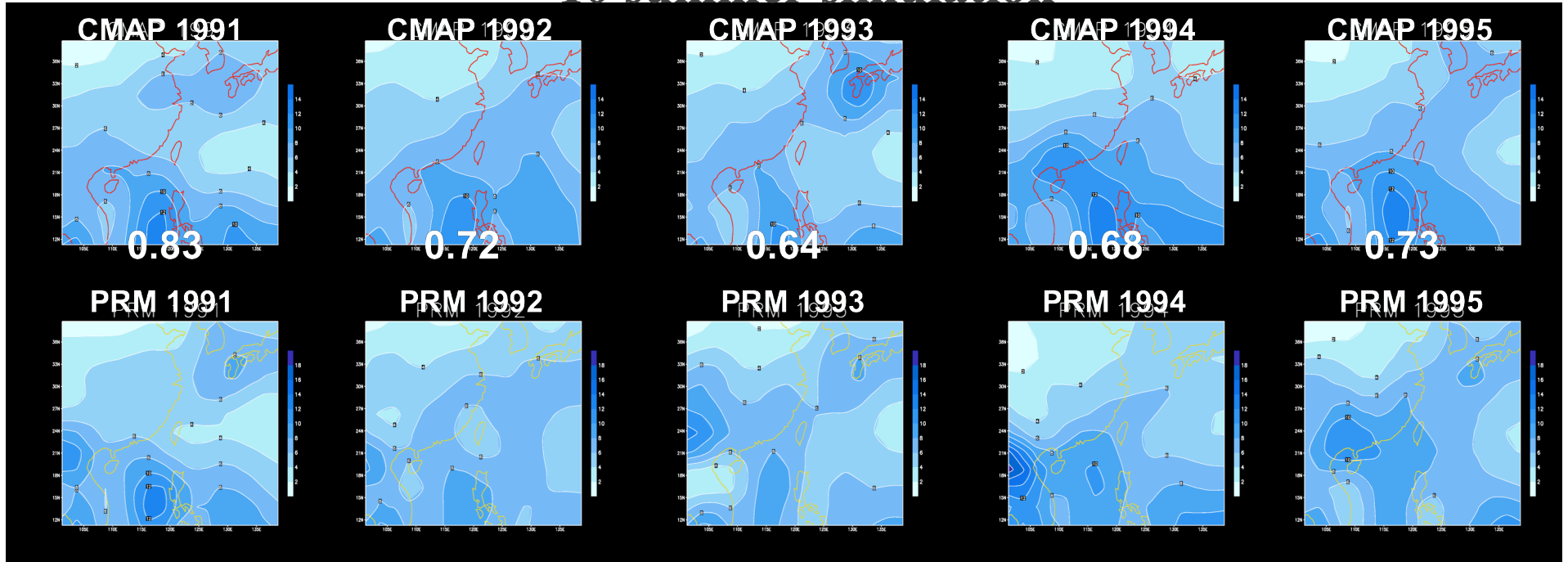
PRM 1997

PRM 1998

PRM 1999

PRM 2000

10 summer simulation



CMAP 1996

CMAP 1997

CMAP 1998

CMAP 1999

CMAP 2000

0.81

0.74

0.66

0.88

0.80

PRM 1996

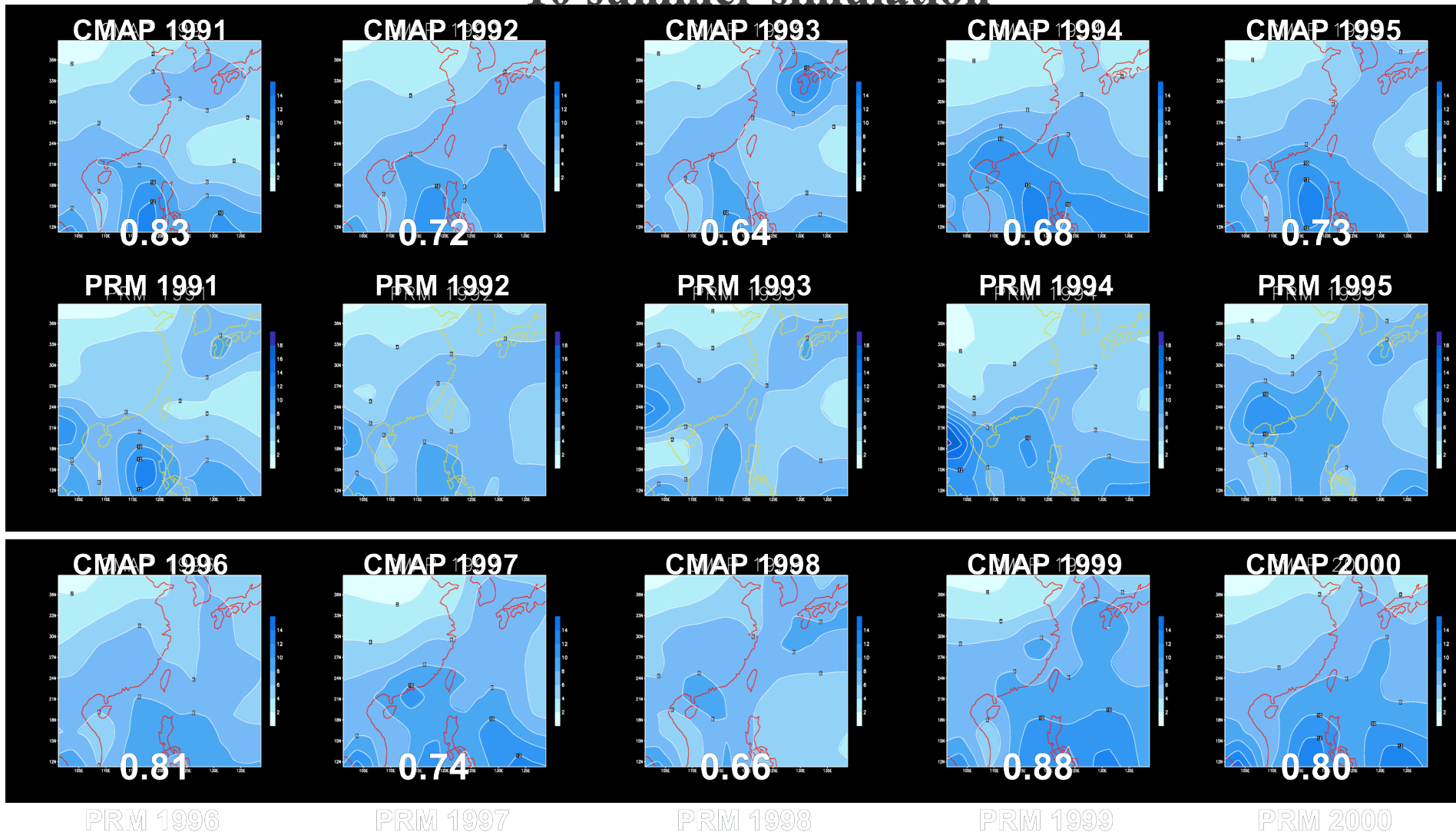
PRM 1997

PRM 1998

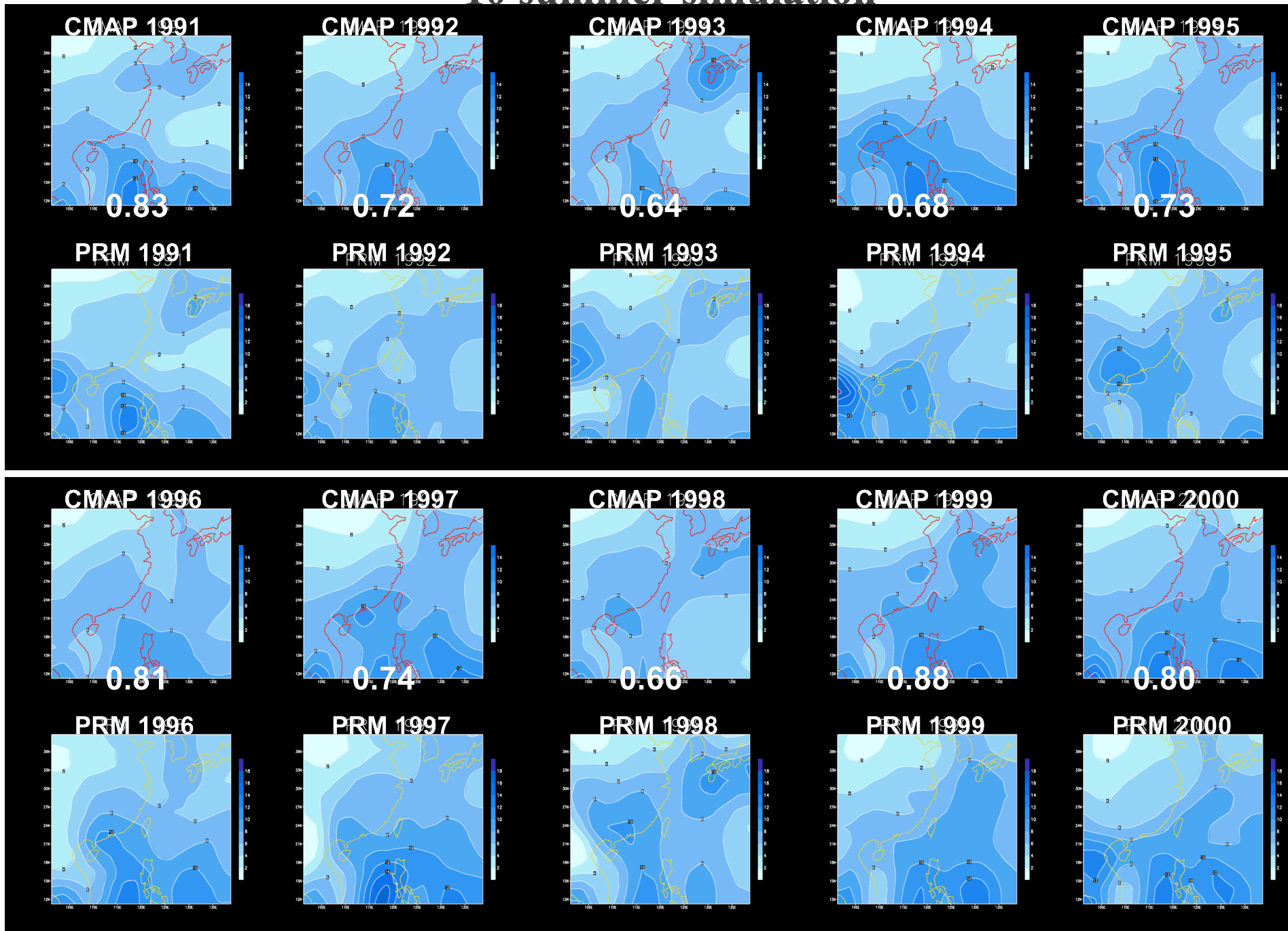
PRM 1999

PRM 2000

10 summer simulation



10 summer simulation

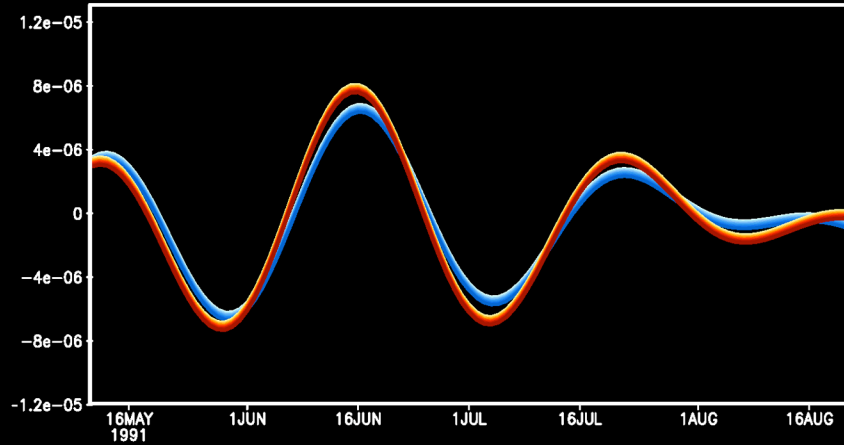


PRM

1991

EC

cor=0.985241

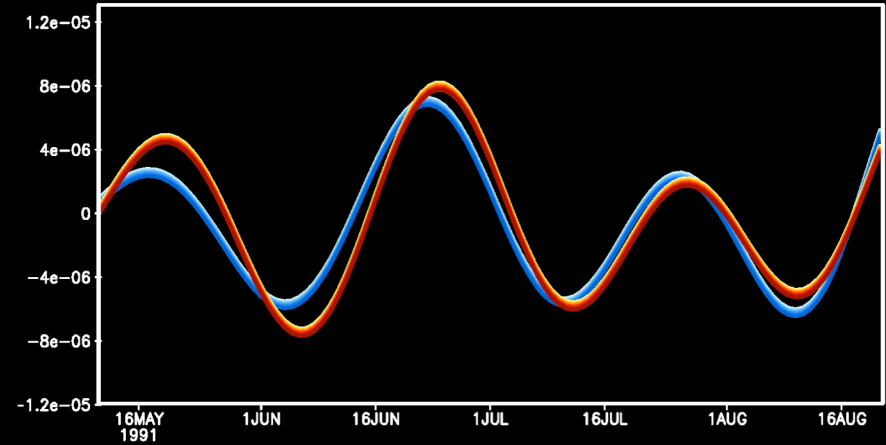


PRM

1992

EC

cor=0.937439



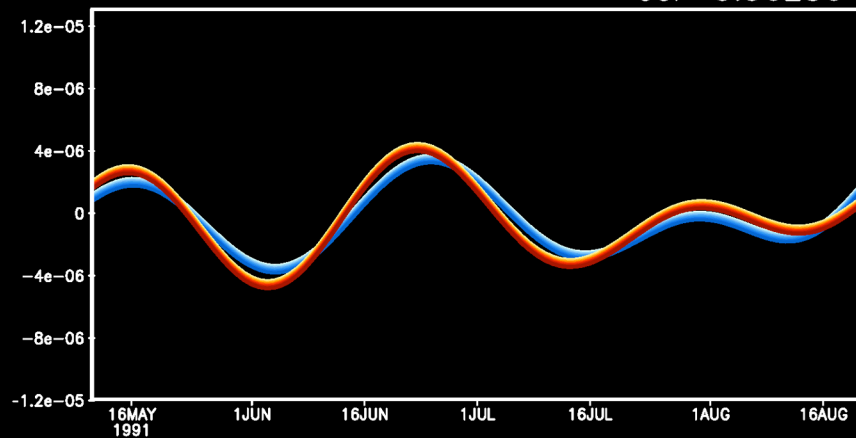
Intraseasonal oscillation in 15~20N, 110~120E

PRM

1993

EC

cor=0.96236

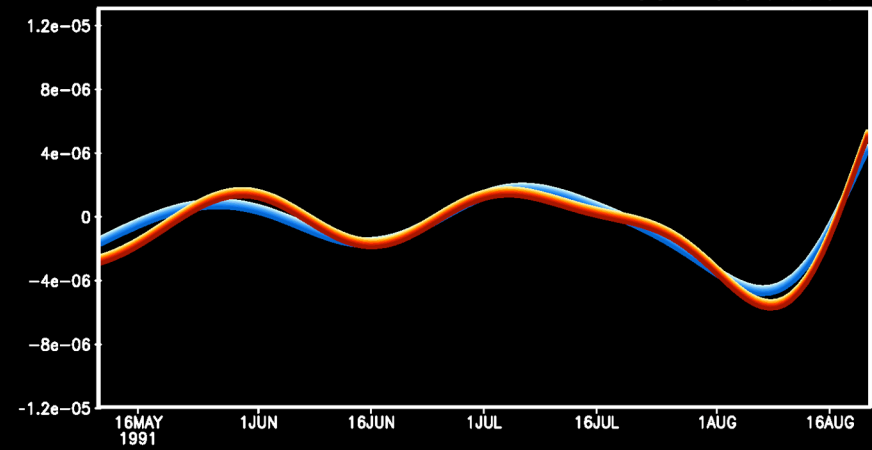


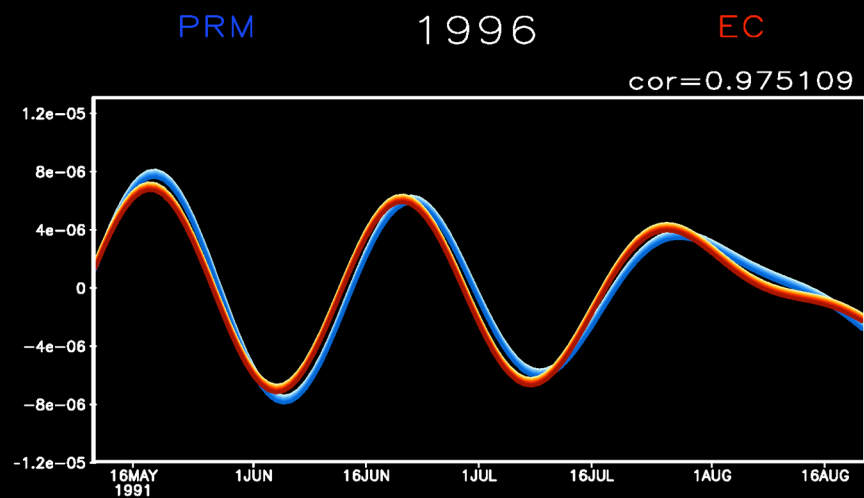
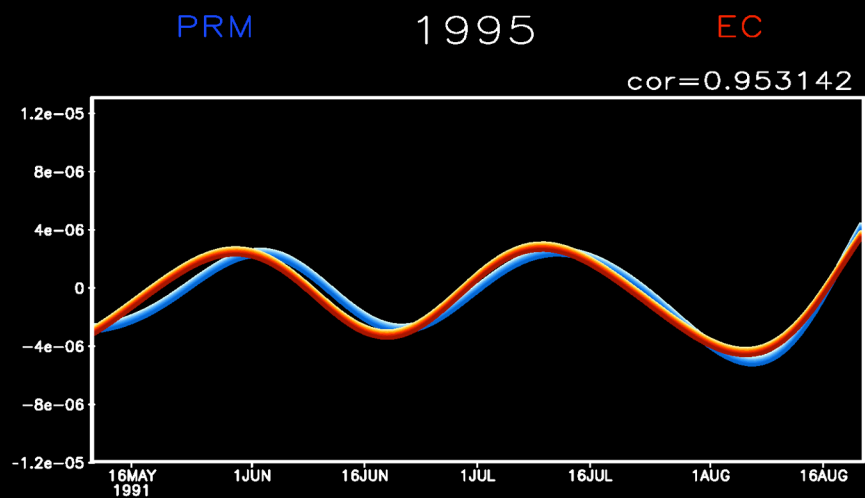
PRM

1994

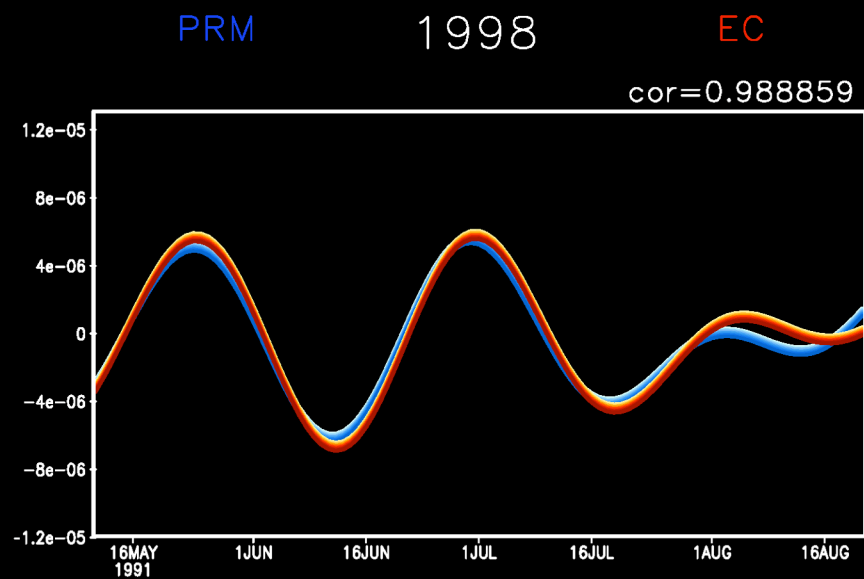
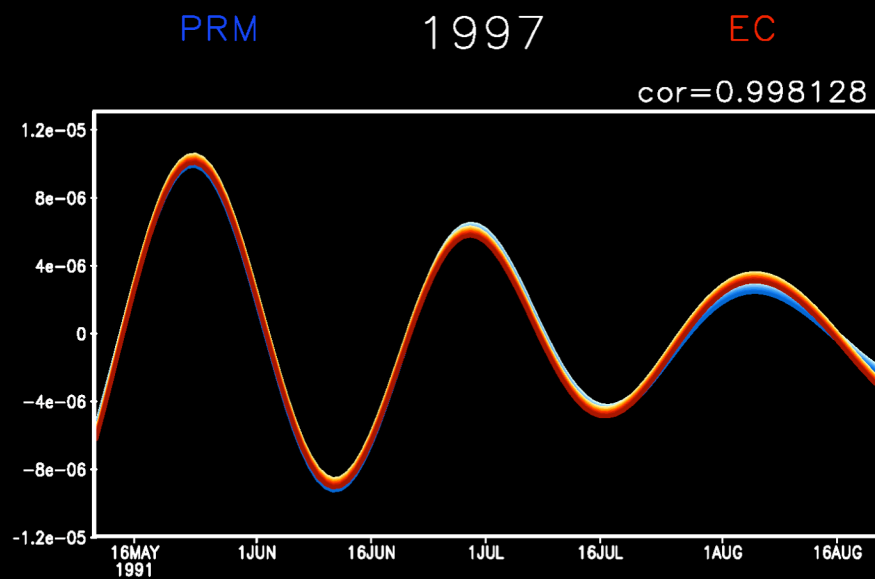
EC

cor=0.964104





Intraseasonal oscillation in 15~20N, 110~120E

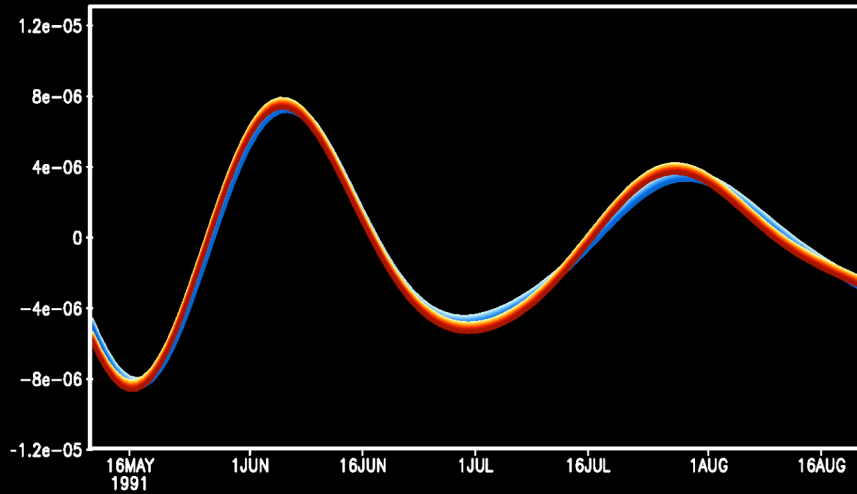


PRM

1999

EC

cor=0.995927

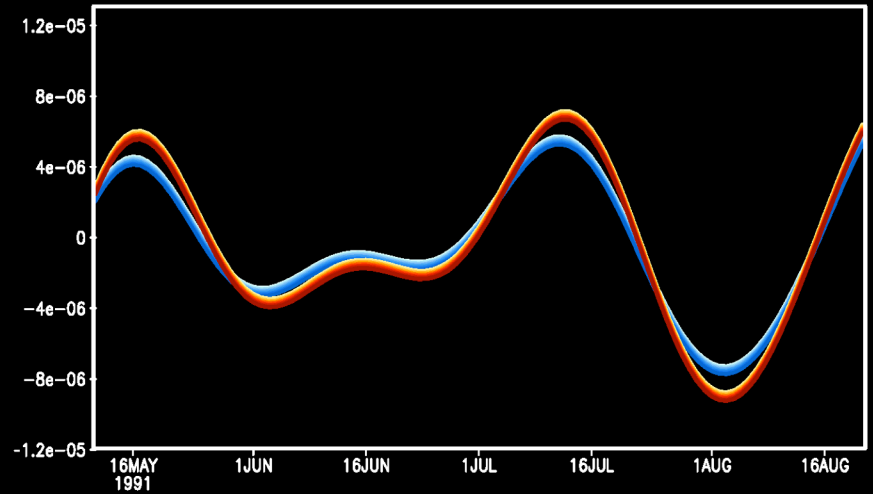


PRM

2000

EC

cor=0.99536



Intraseasonal oscillation in 15~20N, 110~120E

**Thank You for
Your Attention!**

