# Assimilation of GPS RO Data for Typhoon Prediction

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# CASE SIMULATIONS WITH RO DATA:

2001- Typhoon Lekima > 2001- Typhoon Nari > 2002- Typhoon Nakri > 2002- Typhoon Sinlaku > 2003- Dujuan Typhoon (2003/08/31/12-09/02/00) 2004- Cold front-1 (2004/02/06/18) 2004- Cold front-2 (2004/02/07/06) 2004- Meiyu front (2004/05/19/12-22/12) 2004- Conson typhoon (2004/06/07/00-09/00) 2004- Mindulle typhoon (2004/06/29/06-07/02/06) 2004- Aere typhoon (2004/08/21/18-25/12) 2004- Namadol typhoon (2004/12/02-04) 2004- Cold front(2004/12/03 or 06) 2005- Early-March snow event (cold front)\*

# **Model Domain and Physics**

The model simulations use three nested domains at 45-, 15- and 5-km resolutions.

All the simulations use MM5 version 3.5 with explicit treatments (Goddard's scheme) for ice/graupel physics in the three domains (1, 2 and 3), Anthes Kuo's scheme and Grell's scheme for cumulus parameterization in domain 1 (largest) and domain 2, respectively, and the Blackadar scheme for PBL parameterization in all the domains.

3DVAR was performed for each domain with GPSrf.

# **Covariance Matrix-O**

The GPS radio occultation observational covariance matrix is diagonal and thus has assumed no vertical correlations.

This assumption of vertical un-correlation is certainly not supportive of some existing dependence between observations, but in absence of statistical information on those correlations, the assumption insures that the data information is not underestimated in assimilation.

The diagonal elements (variances) are prescribed as a profile exponentially decreasing from 3 N at 100 hPa to 10 N at 1000 hPa. The value of 10 N observational error near the surface is consistent with the 3% refractivity difference between CHAMP radio-occultations and ECMWF analysis found at 1000 hPa, as reported by Kuo et al. (2004).

## Nari best track (CWB)

Such a track was observed once per about thirty years.

Nari made a record-breaking flood in Taipei, causing shutdown of the subway system which took one month to vacate water and clean the facilities, and three months to resume.

The high topography (CMR) in Taiwan plays a major role in rainfall distributions.



# **Observed 24-h accumulated rainfall (mm)**

Major rainfalls over northern, south-western slopes and I-Lan.



(a) 0000-2400 UTC 16 Sep. 2001 (max value: 712 mm)



(b) 0000-2400 UTC 17 Sep. 2001 (max value: 1144 mm)

Wetprf.2001.259.001.01.046.0001.0001.nc	01:45am	51.500	154.030
Wetprf.2001.259.001.01.403.0002.0001.nc	02:22am	38.014	140.001
Wetprf.2001.259.001.01.405.0002.0001.nc	02:26am	25.040	136.964
Wetprf.2001.259.001.01.408.0002.0001.nc	02:30am	4.185	136.568



# 2001/09 Nari Typhoon

9/21

MM5 simulated nearsurface pressure (mb) and wind (ms<sup>-1</sup>) at 24h, 36h.

9/20

NARI,6-19SEP.,2001 typhoon(Vmax>=100kt) typhoon(Vmax 64-99kt) tropical storm(Vmax 34-63kt) tropical depression(Vmax<34kt)

secondray cente

9/18 9/19



> The track is closer to the best track, with GPSrf assimilated.

### 2001 Nari Typhoon

### MM5 simulated 24-h accumulated rainfall (mm)





Some rainfall statistics is closer to the observed, with GPSrf assimilated



Nakri headed for Taiwan from southwest, and was categorized as a tropical storm (not a typhoon yet).



#### Domain-1 (45-km grid)



Domain-1 (45-km)

Wetprf.2002.188.001.01.366.0017.0002.nc Wetprf.2002.188.001.01.367.0017.0002.nc Wetprf.2002.188.001.01.412.0017.0002.nc Wetprf.2002.188.001.01.413.0017.0002.nc Wetprf.2002.188.001.01.417.0017.0002.nc

11:29	22.289	145.435
11:32	34.416	142.594
12:56	5.834	128.987
12:59	16.294	108.563
13:01	22.726	123.296



# 2 4 - h a c c u m u l a t e d rainfall (mm) for the Nakri case.

Threat Score (TS)  
$$TS = \frac{A}{F + O - A}$$

A: the number of the grids on which both forecast and observation exceed the threshold,
F: the number of the grids on which forecast exceeds the threshold, and
O: the number of the grids on which observation exceeds the threshold.

> 1,500 verification grid points on the island.

# **Threat Scores**

Cases	Nari <i>GTS</i>	Nari <i>GPSrf</i>	Nari <i>GTS</i>	Nari <i>GPSrf</i>	Nakri <i>GTS</i>	Nakri <i>GPSrf</i>	Nakri <i>GTS</i>	Nakri <i>GPSrf</i>
Thresholds	0-24 h	0-24 h	24-48 h	24-48 h	0-24 h	0-24 h	24-48 h	24-48 h
0.25 mm	0.545	0.540	0.539	0.536	0.504	0.480	0.505	0.475
0.5 mm	0.544	0.539	0.538	0.535	0.482	0.464	0.498	0.448
1 mm	0.544	0.542	0.532	0.533	0.411	0.425	0.486	0.411
2 mm	0.530	0.524	0.527	0.531	0.376	0.363	0.463	0.378
5 mm	0.489	0.484	0.530	0.528	0.273	0.244	0.321	0.320
10 mm	0.489	0.492	0.527	0.530	0.182	0.174	0.277	0.278
15 mm	0.503	0.496	0.526	0.538	0.149	0.140	0.248	0.258
25 mm	0.524	0.496	0.543	0.531	0.126	0.127	0.233	0.235
50 mm	0.497	0.485	0.462	0.504	0.086	0.088	0.113	0.166
100 mm	0.473	0.444	0.277	0.393	0.009	0.000	0.000	0.129
RMSE (mm)	69.74	71.43	97.84	93.98	90.28	66.55	50.85	37.06

> TS is generally higher for the run with assimilated QuikSCAT data.





Dujuan's simulation, initial time (1200UTC 31 August 2003)  $\pm$  3 hr





Observed rainfall for (a) 00-1200UTC 1 September 2003 (max: 166.806 mm), (b) 12-2400UTC 1 September 2003 (max: 657.744 mm).

# GTS only

### 09/01/00-12UTC

### 09/01/12-24UTC



69.3 mm

369 mm

BOTH (GTS + GPSrf)

### 09/01/00-12UTC

### 09/01/12-24UTC



81.4 mm

387 mm

# **Dujuan Simulation**



# 2004/02 Cold front

#### Initial time:2004/2/6/1800UTC

MM5 simulated nearsurface pressure (mb) and wind (ms<sup>-1</sup>) at 12h, 24h.





## 2004/05 Meiyu Front

#### Initial time:2004/5/19/1200UTC

MM5 simulated nearsurface pressure (mb) and wind (ms<sup>-1</sup>) at 12h, 24h.



# 2004/05 Meiyu Front



MM5 simulated 24-h accumulated rainfall (mm)



**1200UTC 19-20 May 2004** (max value: 162mm )

## 2004/06 Conson Typhoon

a

#### Initial time:2004/6/7/0000UTC

HIVIDAL

h

MM5 simulated nearsurface pressure (mb) and



# 2004/06 Conson

### MM5 simulated 18-h accumulated rainfall (mm)



**0600-2400UTC 8 June 2004** (max value: 76.9mm )



26 1

(a) max value:92.3mm (b) max value:107mm  $\,\circ\,$ 

# 2004/08 Aere Typhoon

Initial time:2004/8/21/1800UTC

2004-08-22\_18:00:00 = 2004-08-21\_18 + 2004-08-22\_18:00:00 = 2004-08-21\_18 +

MM5 simulated nearsurface pressure (mb) and wind (ms<sup>-1</sup>) at 12h,24h, 48h,72h. No AERE.23AUG-2 Acrc(2004/08)--3DVAR (gk) Contour #row 94.48 To 1922.8 CONTOUR INTERVAL OF 2.8888 PT13.3)+ 1871 46 -1 000 BLA PHCS 16D 22.2 CONTOUR INTERVAL OF 2.8888 PT13.3)+ 1871 46 -2.995 BAR UV 2.7 - 1 2284-08-2419.408.00 2284-08-27\_10+72.2841 /25 Aere(2004/08)--3DVAR (gts) 8/24 ARTC(2004)/00/--DIVAR (8(5)) CONTOUR FROM 966.88 TO 1816.8 CONTOUR INTERVAL DF 2.8888 PT(3.3) 1818.40 XFM VECTOR =1.888 5LA PHES IND 1 2884-484-23.154.88.888 = 2884-484-21.154 + 44.884 5MUDIH= 85.1544 -8.995 BARB UV (m/s) 1 2884-884-23.154.88.988 = 2884-484-21.154 + 48.884 5MUDIH= 85.1544 SIGMA SIGMA 8/26 С 8/23 8/22 typhoon(Vmax>=100kt) typhoon(Vmax 64-99kt) tropical storm(Vmax 34-63kt) 10 N tropical depression(Vmax<34kt) 8/20 secondary center 0 0 0 0 0 0 0 0 0 0 10 Aere(2004/08)--3DVAR (gts) DUR FROM 985.00 TO 1024.0 Aere(2004/08)--3DVAR (gts) CONTOUR FROM 976.88 TO 1822.8 CONTOUR INTERVAL OF 2,8888 PT(3,3)= 1811 B.

# 2004/08 Aere Typhoon

MM5 simulated nearsurface pressure (mb) and wind (ms<sup>-1</sup>) at 12h,24h, 48h,72h.

/25

8/26

8/24

8/23

8/22



Acre(2004/08)--3DVAR (gtsrf) CONTOUR FROM 975.80 TO 1822.8 CONTOUR INTERVAL DF 2.8880 PT[3,3]= 1811 MAX



Aere(2004/08)--3DVAR (gtsrf) our from 984.80 to 1824.8 contour interval of 2.88880 pt(3,3)= 1812.43

#### 2004/08 Aere





### 2004/12 Nanmadol Typhoon

MM5 simulated nearsurface pressure (mb) and wind (ms<sup>-1</sup>) at 30h,36h,





### 2004/12 Nanmadol



1200UTC 3-4 December 2004 (max value:528.9mm )



(a) max value:561mm (b) max value:554mm •

#### → About 5-10 % differences

		_	—initial —	- retrieva

The model local refractivity (solid line) and Abelretrieved refractivity (dash line) obtained by ray-tracing near Taipei using the 12-h forecast results at (a) domain 1 (45-km) and (b) domain 2 (15-km) for the Nari case.

$$\frac{d^2 \bar{x}}{d\tau^2} = n \vec{\nabla} n$$
$$(n = 1 + \nu)$$

 $S = \int v d\ell$ 

Sokolovskiy et al. (2004) is an alternative method for excess phase. **Table.** Average relative refractivity (N) error and relative excess phase (S) error below 90 km calculated by nonlocal straightline operator (NSO) for a profile of Abel refractivity.

INT=1: Linear interpolation; INT=2: Cubic interpolation IAB=1: Abel inversion scheme 1; IAB= 2: Abel inversion scheme 2

$\Delta z=0.1 \text{ km}$	IAB=1 (err1, err2)	IAB=2 (err1, err2)
INT=1	(3.117E-04, 6.371E-04)	(1.005E-04, 2.973E-
INT=2	(3.197E-04, 6.562E-04)	(9.826E-05, 2.939E-
$\triangle z=0.2 \text{ km}$	IAB=1 (err1, err2)	<b>143</b> =2 (err1, err2)
INT=1	(6.850E-04, 1.262E-03)	(1.500E-04, 3.360E-
INT=2	(7.171E-04, 1.329E-03)	(1.304E-04, 3.017E-
$\triangle z=0.5 \text{ km}$	IAB=1 (err1, err2)	<b>143</b> =2 (err1, err2)
INT=1	(2.268E-03, 4.257E-03)	(7.525E-04, 1.540E-
INT=2	(2.464E-03, 4.623E-03)	(3.579E-04, 1.152E-
		03)

err1= (N\_mapping – N\_Abel)/N\_Abel from Abel inversion of S err2= (S\_mapping – S\_Abel)/S\_Abel from NSO for N\_mapping



without a spike, (a) dz= 0.1 km (IAB=1, INI=1), err1, err2; (b) dz= 0.5 km (IAB=1, INI=1), err1, err2.

with a spike of N (=130) at z= 10 km, (a) dz= 0.2 km (IAB=1, INI=1), err1 (b) dz= 0.2 km (IAB=2, INI=1), err1, (c) dz= 0.2 km (IAB=1, INI=2), err1.

# Comments on Nonlocal Refractivity Operator Assimilation

- Errors increase considerably near upper boundary, due to the less cancellation effects in a shorter path. Thus, local refractivity operator may still be recommended above the tropopause.
- The refractivity mapping normally has even larger errors compared to errors for excess phases, due to the application of Abel inversion in a finite domain with nonnegligible local refractivity.
- The straightline forward operator is very fast and the assimilation should be computationally much more efficient than bending angle ray-tracing assimilation.

# Comments on Ray-tracing Assimilation for Regional Modeling

Convergence rate in cost-function minimization may be quite low, due to the nonlinearity of the ray path and the sensitivities to any tiny variations of the model state.

The ray may still have considerable bending near the model lateral boundary.

The ray-tracing assimilation is computationally much more expensive than local-refractivity and nonlocal excess phase assimilations.